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# Relationship of ground-level ozone with synoptic weather conditions in Chicago



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#### ABSTRACT

This study investigates the relationship between ground-level ozone variability and synoptic weather conditions in the Chicago area in the summertime over the period 1990-2014. It shows that the occurrence of dry tropical (DT) weather conditions is most likely to lead to high ozone concentrations. Both the 95th percentile of summertime daily maximum 8-hour-average ozone concentrations and the number of days >70 ppb are shown to have a greater and more significant correlation with DT weather in recent years (2004-2014) than in the preceding period (1990-2003), indicating an increased dependence of ozone on DT weather. The DT weather is shown to be associated with increased geopotential height, warmer temperature, and slower eastward wind at 500 hPa and also with warmer temperature and increased northward wind near the surface. The results have implications for the likely effect of future climate change on ozone as a result of modified synoptic weather conditions. More frequent episodes of high ozone concentrations can be expected, leading to worsened effects on public health and the environment. It will also make it increasingly challenging for Chicago to attain the National Ambient Air Quality Standard for ozone, which has recently been tightened by the U.S. Environmental Protection Agency.

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#### 1. Introduction

Ground-level ozone  $(O_3)$  is an important air pollutant. Breathing  $O_3$  can trigger a variety of health problems, particularly for children, the elderly, and people of all ages who have lung diseases such as asthma (Bell et al., 2004; Bell et al., 2007; Lippmann, 1993). It can also have harmful effects on sensitive vegetation and ecosystems (Fishman et al., 2010; Heagle, 1989; Kline et al., 2008). Ground-level  $O_3$  is formed through photochemical oxidation of carbon monoxide and volatile organic compounds (VOC) by the hydroxyl radical (OH) in the presence of nitrogen oxides (NO<sub>x</sub>). Ozone is removed from the atmosphere by its chemical reaction with NO<sub>x</sub> and its photolysis in the presence of water vapor, as well as by dry deposition (Jacobson, 2002). The production of ground-level  $O_3$  is affected not only by the emissions of its precursors (i.e., NO<sub>x</sub> and VOC), but also by meteorological conditions. Ozone production is greater on warm, sunny days when the air is stagnant than on days when it is cool, cloudy, and windy (Cox and Chu, 1996; Lee et al., 2012). The influences of meteorology on  $O_3$  include the transport of  $O_3$  and its precursors (i.e., NO<sub>x</sub> and VOC) by air flow, as well as the temperature and amount of sunlight, which both affect the chemical reaction rates of  $O_3$  production and destruction.

The U.S. Clean Air Act of 1970 requires the U.S. Environmental Protection Agency (US EPA) to set National Ambient Air Quality Standards (NAAQS) for  $O_3$  and other pollutants that are considered harmful to public health and the environment (so-called "criteria" pollutants). The primary NAAQS for  $O_3$  (to protect public health) and the secondary NAAQS (to protect public welfare) had both been set at the level of 75 ppb, measured as the annual fourth–highest daily maximum 8-hr concentration, averaged over three years. After an extensive review of the health effects of  $O_3$  by the Clean Air Scientific Advisory Committee, the US EPA tight-ened the ambient standard to 70 ppb in October 2015, with the intention to extend the health protection to people at higher risk of  $O_3$  pollution (e.g., people with respiratory conditions) and sensitive ecosystems (US EPA, 2015).

Ozone levels are usually limited by the supply of NO<sub>x</sub>, but O<sub>3</sub> can also be VOC-limited in highly-polluted urban areas (Jacob and Winner, 2009). Ozone control strategies include reducing VOC and NO<sub>x</sub> emissions. NO<sub>x</sub> emissions come mainly from combustion activities, while VOC have both anthropogenic and biogenic sources (Jacobson, 2002). Efforts in the U.S. since 1998 have focused on controlling NO<sub>x</sub> emissions. Because of the effects of regulatory actions since the passage of the Clean Air Act and its 1977 Amendments, national average O<sub>3</sub> concentrations in the U.S. decreased by 22% from 1983 to 2002. The number of high surface O<sub>3</sub> events was also reduced in most cities during this period (US EPA, 2003). Both of these achievements reflect the success of efforts made by the US EPA and state and tribal agencies to drive down emissions of the O<sub>3</sub> precursors, NO<sub>x</sub> and VOC. Nevertheless, the rate of improvement in O<sub>3</sub> levels slowed in the recent decade (a decrease of only 13% in national average O<sub>3</sub> concentrations from 2000 to 2010), and many metropolitan areas of the U.S. continue to violate the NAAQS for O<sub>3</sub>, including Chicago (US EPA, 2012). The diminishing returns of emission reductions on O<sub>3</sub> concentrations at least partly reflect the growing importance of weather.

It has been known for many years that ground-level  $O_3$  is highly dependent on weather (see, e.g., Bloomfield et al., 1996; Camalier et al., 2007; Cox and Chu, 1996; Lee et al., 2012; Shen et al., 2015). Several studies have attempted to quantify the effect of meteorological variability on  $O_3$  concentrations, with values such as 53% for the Chicago area (Jing et al., 2014), 56–80% for 39 eastern U.S. urban areas (Camalier et al., 2007), and 50–80% for 74 areas across the eastern U.S. (Davis et al., 2011). Because of this sensitivity to weather, it follows that climate change should influence  $O_3$  (Bell et al., 2007; Holloway et al., 2008; Jacob and Winner, 2009; Oswald et al., 2015). Warmer temperatures can be expected to lead to higher  $O_3$  concentrations due to a greater  $O_3$  production rate, as well as to increased production of biogenic emissions of VOC and  $NO_x$  (Lin et al., 2007). In addition, anthropogenic emissions of  $NO_x$  in summertime are expected to increase in a warmer climate due to greater energy consumption for summertime cooling (Hadley et al., 2006; Lin et al., 2010; McDonald et al., 2012).

Climate change could also affect  $O_3$  through modifications of synoptic-scale weather systems. It has been shown that climate change has resulted in decreased frequency of mid-latitude cyclones (Key and Chan, 1999; Knapp and Soulé, 2007). It has also been found that summertime  $O_3$  concentrations are negatively correlated with mid-latitude cyclone frequencies; the decreased frequency of mid-latitude cyclones between 1980 and 2006, which resulted in more stagnant weather, largely offset the improvement in  $O_3$  in the eastern U.S. that was obtained by reducing anthropogenic emissions (Leibensperger et al., 2008). It has also been predicted that cyclone activities will continue to decrease in the future due to climate change (Turner et al., 2013) Download English Version:

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