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Short Communication

Permeability measurements of quartz sands with methane hydrate

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ABSTRACT

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HIGHLIGHTS

• We measured water effective permeability of sands with/without hydrate.

• Shape factor in hydrate-bearing sediments was associated with hydrate saturation.

• Ratio of permeability with/without hydrate was related to the hydrate saturation.

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1. Introduction

Considered to be a potential energy resource, Natural Gas Hydrates (NGH) formed by water and gas molecules are widely distributed in marine and permafrost sediments (Collett et al., 2015). There have been no field tests developed that are capable of indicating the commercial recovery of gas from hydrate reservoirs under reliable permeability estimation. The permeability of porous media with hydrate is one of the most critical parameters that determines the gas recovery potential from natural gas hydrate reservoirs (Li et al., 2010).

Permeability measurements in hydrate-free and/or hydratebearing sediments have been reported. Minagawa et al. (2008) measured the permeability of artificial sands and Mallik simulated sands with and without methane hydrates, and compared the results using Darcy's Law and the NMR spectra method. Kumar et al. (2010) measured the permeability of glass beads with and without CO₂ hydrates, and the permeability without hydrates was 66.9 Darcy. Sakamoto et al. (2010) measured the water permeability with methane hydrates in three types of sands, and the results indicated that the parameter was sensitive to the types of porous media, the hydrate-forming gas and the hydrate saturations.

The permeability of porous media is one of the critical parameters that determines gas recovery from nat-

ural gas hydrate reservoirs. We measured the permeability of quartz sands by injecting water at a certain

flow rate. In hydrate-bearing sands, the larger the hydrate saturation, the smaller the shape factor in the

presence of hydrate. A k_r - S_H relationship with the hydrate saturation both lower and higher than 10%

was obtained. Based on the experimental results, the saturation exponent *n* varied from 6.0 to 1.0.

In this study, we measured the permeability of quartz sands with and without methane hydrates by injecting deionized water. A k_r - S_H relationship with an S_H both lower and higher than 10% was obtained.

2. Experiments

The complete descriptions of the apparatus and the materials have been introduced by Li et al. (2017). Table 1 shows the properties of the quartz sands. The grain size (the volume weighted mean diameter) of the guartz sands were measured using the Mastersizer 2000E (0.0001-1.0000 mm) produced by Malvern Instruments Ltd, United Kingdom. The effective permeability $k (m^2)$ was calculated using Darcy's Law:



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Table 1
Properties of the quartz sands

Parameter	Quartz Sands I (QS I)	Quartz Sands II (QS II)	Quartz Sands III (QS III)	Quartz Sands IV (QS IV)
Sieve mesh	30-40	40-60	60-80	80-120
Grain size (volume weight mean diameter, mm)	0.7085	0.5502	0.3332	0.2316
Density of quartz sands (kg/m ³)	2603.0	2599.1	2593.4	2615.7
Mass of quartz sands in the vessel (g)	120.05	118.31	116.28	117.98
Total pore volume (mL)	37.52	39.28	38.59	41.72
Porosity (%)	45.87	48.02	47.18	51.00
Volume weighted mean diameter (µm)	731.9	572.0	336.1	226.6
Specific surface area (m ² /g)	0.038	0.053	0.114	0.087

$$k = \frac{\mu QL}{A\Delta P} \tag{1}$$

where μ is the dynamic viscosity of the fluid (Pa s). *Q* is the fluid flow rate (m³/s). *L* = 0.250 m and *A* are the inner length and the inner cross-sectional area (with the radius of 0.010 m) of the cylindrical vessel. ΔP is the differential pressure (Pa). 1 Darcy = 0.987 × 10⁻¹² m². The absolute permeability k_0 was considered to be equal to the water effective permeability k_w , which was measured by injecting water with a certain flow rate under stable pressure and temperature.

Table 2 shows the experimental conditions and the results of the permeability measurements in the hydrate-free sands. For k_0 of QS I, there was no statistical difference between Runs 1–8 and 9–16 under different system pressures. This confirmed the results of the previous study by Li et al. (2013). From the coarse sands of QS I to the fine sands of QS IV, k_0 decreased significantly from over 50.0 Darcy to approximately 14.0 Darcy. The main reason for this was that with the reduction of the particle size, the flow channels in the porous media became narrow and the flow resistance increased significantly.

The hydrate formation process and the calculations of the gas (S_G) and hydrate (S_H) saturations have been described by Li et al. (2017). Table 3 shows the experimental conditions and the results of the permeability measurements in the hydrate-free and hydrate-bearing sands. For the system with low gas saturation

(<30%), the gas could be driven out by the injected water quickly and effectively. The permeability in hydrate-bearing sands $k(S_H)$ could be calculated using Eq. (1).

3. Results and discussion

3.1. Kozeny–Carman equation

Predicting the absolute permeability of porous media is difficult, especially for particles with widely distributed grain sizes. The Kozeny–Carman equation (Kozeny, 1927; Carman, 1937) is a widely used model. It successfully relates the permeability of porous media with the parameters that describe the properties of the pores. A simple form (Hearst et al., 2000; Kleinberg et al., 2003) of the Kozeny–Carman equation refers to the internal surface area of the pore space and the total pore volume:

$$k_0 = \frac{\phi}{v\tau (A_{pore}/V_{pore})^2} \tag{2}$$

where k_0 is the absolute permeability (Darcy). ϕ is the porosity (%). v is the shape factor. τ is the tortuosity. A_{pore} is the internal pore surface area (m²). V_{pore} is the total pore volume (m³).

In a previous study by Kleinberg et al. (2003), the shape factor v in the Kozeny–Carman equation was on the order of unity. However, the micro-structure of the grains and the pore spaces may

Table 2

Experimental conditions and results of the permeability measurements in the hydrate-free sands.

Parameter	Runs	P (MPa)	T (°C)	μ_w (×10 ⁻³ Pa s)	ΔP (kPa)	Q (×10 ⁻⁷ m^3/s)	k ₀ (Darcy)
QS I	1	0.10	8.42	1.368	3.16	1.677	58.5
	2	0.10	8.55	1.363	12.43	4.995	44.2
	3	0.10	8.33	1.372	13.12	4.912	41.4
	4	0.10	14.79	1.144	5.41	3.328	56.8
	5	0.10	8.47	1.366	7.75	3.248	46.2
	6	0.10	8.46	1.367	4.54	2.497	60.6
	7	0.10	8.37	1.370	3.76	1.614	47.5
	8	0.10	14.73	1.146	3.30	1.603	44.9
	9	15.48	8.52	1.347	9.87	4.939	54.4
	10	15.28	8.39	1.352	11.18	4.916	48.0
	11	15.43	8.47	1.349	8.19	4.083	54.2
	12	15.56	8.39	1.352	6.53	3.280	54.7
	13	15.68	8.38	1.352	7.09	3.267	50.2
	14	15.39	8.39	1.352	4.77	2.436	55.7
	15	15.42	8.52	1.347	3.15	1.607	55.3
	16	15.45	8.53	1.346	3.41	1.448	46.1
QS II	II-1	0.10	8.50	1.365	4.68	1.511	35.5
QS III	III-1	0.10	8.34	1.371	8.78	1.675	21.1
QS IV	IV-1	15.32	8.50	1.348	12.36	1.646	14.5
	IV-2	15.35	8.64	1.342	6.23	0.832	14.4
	IV-3	15.20	8.38	1.353	6.29	0.795	13.8
	IV-4	15.07	8.57	1.345	6.19	0.805	14.1
	IV-5	15.28	8.43	1.350	12.02	1.612	14.6
	IV-6	15.31	8.40	1.352	12.11	1.575	14.2
	IV-7	15.06	8.54	1.346	6.24	0.831	14.5
	IV-8	15.46	8.53	1.346	12.19	1.649	14.7
	IV-9	15.08	8.52	1.347	6.37	0.831	14.2

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