

Short communication

Lead-210 and Beryllium-7 fallout rates on the southeastern coast of Brazil

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ABSTRACT

Total ²¹⁰Pb and ⁷Be fallout rates were measured on the coastal region of Niteroi, Brazil. The monthly depositional flux of ²¹⁰Pb and ⁷Be varied by a factor of 26, from 1.7 to 43.3 mBq cm⁻² year⁻¹ and ~27, from 7.5 to 203.5 mBq cm⁻² year⁻¹, respectively. The relatively large oscillations in the depositional flux of ²¹⁰Pb at this study site were likely due to variations in air mass sources, while the ⁷Be fluctuations may be driven by a combination of weather conditions. Local geology could support the periodic high fluxes of ²¹⁰Pb from continental air masses, as shifting oceanic wind sources were affirmed by the uncorrelated ²¹⁰Pb and ⁷Be fallout activities and ⁷Be/²¹⁰Pb ratios. The ²¹⁰Pb atmospheric deposition was found to be in agreement with local sediment inventories, an important consideration in geochemical studies that estimate sedimentation processes.

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1. Introduction

Radium-226 within rocks and sediments produces ²²²Rn as part of the ²³⁸U decay series. Radon-222, an inert gas, diffuses into the atmosphere and decays through several short lived daughter nuclides to ²¹⁰Pb (*t*_{1/2}22.3y). Beryllium-7 (*t*_{1/2}53d) is a relatively short lived naturally occurring radionuclide of cosmogenic origin. It is formed in the upper atmosphere by spallation of light atmospheric nuclei, nitrogen (Z ¼ 7), oxygen (Z ¼ 8) and carbon (Z ¼ 6) when they interact with cosmic rays.

Since ⁷Be is of atmospheric origin, it is an excellent tracer of atmospheric sources and transport processes. Both ²¹⁰Pb and ⁷Be may be removed from the atmosphere via wet or dry deposition (Papastefanou and Ioannidou, 2004; Kulan et al., 2006; Caillet et al., 2001). Information on ²¹⁰Pb and ⁷Be-derived fallout rates of aerosols is useful to determine the transport processes of other atmospheric contaminants that behave in a similar manner to these radionuclides, such as trace metals (Kim et al 2000; McNeary and Baskaran, 2003). Seasonal events, such as rain, temperature

variability and winds influence the distribution of ²¹⁰Pb and ⁷Be fallout in depositional areas (Al-Azmi et al., 2001; Leppanen et al., 2010).

Recent ⁷Be atmospheric studies in Rio de Janeiro show that regional precipitation and irregular air-masses patterns influence the ⁷Be variability in the near surface air concentrations (Pacini et al., 2011; Leppanen et al., 2010). In the present study, we report on variations in the total atmospheric fluxes of ²¹⁰Pb and ⁷Be to Niteroi – Rio de Janeiro, Brazil. This study also compares ²¹⁰Pb annual flux rates in local sediment columns to the ²¹⁰Pb fallout rates determined in this work. Knowledge of regional ²¹⁰Pb and ⁷Be fallout rates are particularly important in geochemical studies that determine depositional processes such as sediment mixing and accumulation (e.g., Smoak and Patchineelam, 1999; Sanders et al., 2010a; Sabaris and Bonotto, 2011).

The maximum rainfall in the state of Rio de Janeiro occurs between December and March and minimum between June and August. On a monthly basis, rough estimates from local meteorological station indicate prevailing south-southeastern winds between April and September and north-northeastern winds from October to May. The coastal regional geology in southeastern Brazil is lined by the Serra do Mar mountain range and dominated by igneous rocks such as granite and granodiorite.

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2. Methods

Using a modified approach of Schuler et al. (1991), a total deposition collector consisting of a 20-L polyethylene drum with surface area of 2500 cm² was setup in Niterói, RJ, Brazil (22°53.55'S, 43°08.95'W) on a stand, near mean sea level, to avoid contamination by resuspended dust (Fig. 1). The collector was continuously exposed to the atmosphere from June 2006 through May 2007 and sampled monthly (Table 1).

To prevent any potential adsorption of ⁷Be and ²¹⁰Pb on to the walls of the collector, the container was acidified prior to deployment with concentrated HCl. Total deposition samples were collected roughly every month or when drum was nearly full. Immediately after collection, yield spikes [1 mL ICP- Standard solution (10 mg L⁻¹ Be, 10 mg L⁻¹ Pb)] were added to the sample. The drum was cleaned repeatedly with 6 M HCl to remove any adsorbed Be and Pb from the surface of the drum. Nine hundred mg of iron, in the form of Fe(NH₄)₂(SO₄)₂, was added to samples and allowed to equilibrate for 3 h. Lead and beryllium were coprecipitated with iron by adding NH₃OH to attain a pH of ~7. After a minimum of 24 h, the Fe(OH)₃ coprecipitate was collected by decanting the supernatant and centrifuging (modification of Schuler et al., 1991). The solution in the centrifuge flask was then dried and quantitatively transferred to a gamma counting vial.

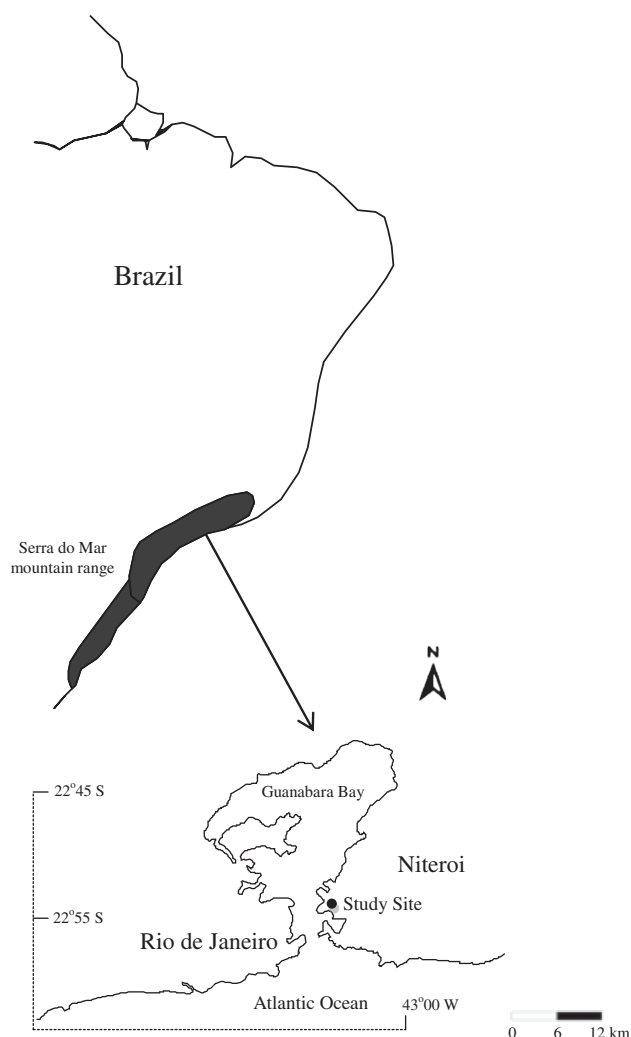


Fig. 1. Study site (Niterói, Brazil).

Table 1

Sampling period (12 months) from June 1st 2006 to May 31st 2007, average sea level monthly temperatures (°C), rainfall (mm), ²¹⁰Pb, ⁷Be and ⁷Be/²¹⁰Pb (mBq cm⁻² year⁻¹).

Sampling	Temp.	Rainfall	²¹⁰ Pb	⁷ Be	⁷ Be/ ²¹⁰ Pb
Jun	22	35	43.7	51.7	1.2
Jul	22	60	11.0	32.0	2.9
Aug	23	34	4.6	37.9	8.3
Sep	22	90	17.2	63.5	3.7
Oct	23	97	1.9	7.5	4.0
Nov	23	116	29.5	203.5	6.9
Dec	26	95	39.1	61.9	1.6
Jan	26	140	4.7	40.6	8.7
Feb	27	76	9.4	9.1	1.0
Mar	27	11	12.8	28.8	2.3
Apr	26	47	6.5	33.9	5.2
May	22	64	2.9	28.8	10.1
Yearly mean	24	72	15.3	49.8	4.6

Gamma counting was conducted in a high purity germanium coaxial high purity well coaxial detector, housed in a lead shield and coupled to a multichannel analyzer. Lead-210 and ⁷Be were quantified by integrating gamma peaks at 46.5 and 477.6 keV respectively. Activities were calculated by multiplying the counts per minute by a factor that includes the gamma-ray intensity and detector efficiency. The ²¹⁰Pb counting efficiency was determined directly by the counting of a standard material. Due to its short half a ⁷Be standard was not available. The ⁷Be counting efficiency was interpolated from an exponential curve fit ($r^2 = 0.98$) of several detector efficiencies over an energy range from 46.5 to 911 keV. Counting errors were 2–3% for ⁷Be and less than 5% for ²¹⁰Pb. Self absorption was evaluated by reducing the mass of a sample and repeating analysis. The sample mass was reduced seven times over a mass range from 1.7 to 0.08 g. The activity (2016.6 ± 90 mBqg⁻¹) remained the same with each reduction in mass. The net disintegrations per minute (dpm) vs. mass showed a linear relationship indicating no self adsorption occurred within this mass range.

The overall chemical yield of this procedure was determined after gamma analyses were completed. The sample was dissolved in 100 mL of 3 M HCl and 5 mL of this solution was analyzed for stable Be and Pb using an atomic absorption spectrometer (AAS). The chemical yield was used in the final calculation of ²¹⁰Pb and ⁷Be.

3. Results and discussion

Collection intervals with specific fallout rates of ²¹⁰Pb and ⁷Be for the 12 month period studied are shown in Table 1. The total ⁷Be and ²¹⁰Pb depositional fluxes in the study site fluctuated over an order of magnitude during the sampling period. The ²¹⁰Pb fallout rate varied by a factor of 26, from 1.7 to 43.3 mBq cm⁻² year⁻¹. The corresponding fallout rate of ⁷Be varied by a factor of ~27, from 7.5 to 203.5 mBq cm⁻² year⁻¹. The oscillations in the depositional fluxes do not appear to be influenced by rainwater flow or temperature fluctuations (Fig. 2). Indeed, no significant correlations were found between specific activities and the climatic variables shown in Fig. 2 or general wind direction (continental or oceanic).

Fallout rates depend on air mass source and a wide variation in air mass source can produce a relatively large variation in fallout rates. Due to low exhalation rates of ²²²Rn over the ocean as compared to continental areas, oceanic air masses are typically depleted in ²²²Rn and its daughter products, including ²¹⁰Pb (Baskaran and Swarzenski, 2007). As a consequence, atmospheric ²¹⁰Pb has a strong longitudinal dependency attributed to air mass origins (Turekian et al., 1977). However, ⁷Be is of cosmogenic origin, and its flux to the Earth's surface has a latitudinal dependence,

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