

## Technical Communication

## Axisymmetric consolidation of a poroelastic soil layer with a compressible fluid constituent due to groundwater drawdown

Kang-He Xie<sup>a</sup>, Da-Zhong Huang<sup>a,\*</sup>, Yu-Lin Wang<sup>b</sup>, Yue-Bao Deng<sup>c</sup><sup>a</sup> Research Center of Coastal and Urban Geotechnical Engineering, Zhejiang University, Hangzhou 310058, China<sup>b</sup> Department of Environment and Civil Engineering, Wuyi University, Wuyi Shan 354300, China<sup>c</sup> Institute of Geotechnical Engineering, Ningbo University, Ningbo 315211, China

## ARTICLE INFO

## Article history:

Received 30 June 2013

Received in revised form 19 October 2013

Accepted 20 October 2013

Available online 9 November 2013

## Keywords:

Axisymmetric consolidation

Groundwater drawdown

Compressibility of fluid

Laplace–Hankel transform

Analytical solution

## ABSTRACT

Axisymmetric consolidation of a poroelastic soil layer with a compressible fluid constituent induced by groundwater drawdown was studied based on Biot's axisymmetric consolidation theory. Laplace and Hankel transforms were employed to solve the governing equation. Explicit analytical solutions are obtained in the Laplace–Hankel transform domain when groundwater drawdown is induced by a constant pumping well. Based on the solutions, numerical computations were performed to study the influences of the compressibility of the fluid constituent on the consolidation behavior of the soil layer.

© 2013 Elsevier Ltd. All rights reserved.

## 1. Introduction

Land subsidence is a worldwide problem that has caused substantial damage to infrastructure and resulted in great economic losses [1,2]. The problem is more severe in regions with soft soils where groundwater is heavily exploited for industrial and domestic purposes [3]. Two large regions in China that suffer from severe subsidence problems are the Yangtze River Delta and the North China Plain, where thick, soft soil layers exist [3–6]. In each of these regions, as groundwater is pumped from the aquifer, the water level and pore water pressure are reduced, which leads to an increase in effective stress and consolidation of the soil layer and thus results in land subsidence.

Because the permeability of an aquifer is usually much greater than that of a soil layer, the consolidation of a soil layer can be analyzed in two steps [7]. The first step is to compute or measure the groundwater drawdown in the aquifer. The second step is to compute the consolidation of the soil layer using the results of the first step as the boundary conditions. Several analytical solutions have been presented for the one-dimensional consolidation of a soil layer due to groundwater drawdown. Luo et al. [8] presented solutions for the one-dimensional consolidation of a soil layer due to constant groundwater drawdown. Li and Helm [9]

presented solutions for the one-dimensional consolidation of a soil layer subjected to periodic groundwater fluctuation. Li [10] studied the one-dimensional consolidation of a nonlinear elastic soil layer subjected to periodic groundwater fluctuation. Teng et al. [11] studied the one-dimensional consolidation of a soil layer caused by constant groundwater drawdown with consideration of the body force effect. Tsai [12] studied the one-dimensional soil consolidation caused by constant groundwater drawdown with consideration of the viscosity effect. Liu et al. [13] presented solutions for the one-dimensional consolidation of a visco-elastic soil layer due to withdrawal of groundwater.

In one-dimensional consolidation theory, groundwater drawdown is assumed to be widely and uniformly distributed in an aquifer. However, because there are different types of groundwater drawdown that correspond to different pumping conditions used in practice, it is necessary to study the consolidation of soil layers based on Biot's three-dimensional consolidation theory [14,15]. On the other hand, the degree of saturation of a soil layer is usually not 100% because the soil may contain small bubbles. It has been reported that even a very small amount of gas in a soil dramatically increases the compressibility of the fluid [16]. Some researchers have studied the importance of fluid compressibility in the consolidation problem [16–20]. However, there have been no studies on the consolidation of a poroelastic soil layer with a compressible fluid constituent due to groundwater drawdown.

\* Corresponding author.

E-mail address: [huangdz05@163.com](mailto:huangdz05@163.com) (D.-Z. Huang).

## Nomenclature

$u_r$	radial displacement	$t$	time variable
$u_z$	vertical displacement	$s$	Laplace transform variable
$r, \theta, z$	cylindrical coordinates	$\xi$	Hankel transform variable
$\sigma_{ij}$	stress component	$J_n$	$n$ th-order Bessel function of the first kind
$p$	excess pore water pressure	$a$	conductivity of the aquifer
$\varepsilon_v$	volumetric strain	$k_2$	coefficient of permeability of the aquifer
$\varepsilon_{ij}$	strain component	$Q$	constant pumping rate
$G$	shear modulus	$H$	thickness of the soil layer
$k$	coefficient of permeability	$H_2$	thickness of the aquifer
$\gamma_w$	unit weight of water		
$M$	bulk modulus of pore fluid		
$\nu$	Poisson's ratio		

This paper presents a study of the axisymmetric consolidation of a poroelastic soil layer with a compressible fluid constituent due to groundwater drawdown based on Biot's axisymmetric consolidation theory. Laplace and Hankel transforms were employed to solve the governing equation. When groundwater is pumped from an aquifer at a constant rate, explicit analytical solutions are obtained in the Laplace–Hankel transform domain. Based on the solutions for groundwater drawdown induced by a constant pumping well, numerical computations were performed to study the influences of the compressibility of the fluid constituent on the consolidation behavior of the soil layer.

## 2. Mathematical model

### 2.1. Governing equations

It is assumed that the poroelastic layer is homogenous isotropic. The force equilibrium equations (with no body forces) specialized to axisymmetry are

$$\frac{\partial \sigma_{rr}}{\partial r} + \frac{\partial \sigma_{rz}}{\partial z} + \frac{\sigma_{rr} - \sigma_{\theta\theta}}{r} = 0 \quad (1)$$

$$\frac{\partial \sigma_{rz}}{\partial r} + \frac{\partial \sigma_{zz}}{\partial z} + \frac{\sigma_{rz}}{r} = 0 \quad (2)$$

where  $\sigma_{rr}$ ,  $\sigma_{\theta\theta}$ ,  $\sigma_{zz}$  are total normal stress components and  $\sigma_{rz}$  is shear stress component.

Assuming the solid constituent is incompressible, the constitutive equations take the form

$$\sigma_{ij} = 2G\left(\varepsilon_{ij} + \frac{\nu}{1-2\nu}\varepsilon_v\delta_{ij}\right) - p\delta_{ij} \quad (3)$$

where the subscripts  $i$  and  $j$  can be  $r$ ,  $\theta$  or  $z$ ,  $G$  is the shear modulus,  $\nu$  is the Poisson's ratio,  $p$  is the excess pore water pressure,  $\delta_{ij}$  is the Kronecker delta.

The Navier-type equations for displacement and  $p$  are obtained by substituting Eq. (3) into Eqs. (1) and (2):

$$\left(\nabla^2 - \frac{1}{r^2}\right)u_r + \frac{1}{1-2\nu}\frac{\partial \varepsilon_v}{\partial r} - \frac{1}{G}\frac{\partial p}{\partial r} = 0 \quad (4)$$

$$\nabla^2 u_z + \frac{1}{1-2\nu}\frac{\partial \varepsilon_v}{\partial z} - \frac{1}{G}\frac{\partial p}{\partial z} = 0 \quad (5)$$

where  $\nabla^2 = \frac{\partial^2}{\partial r^2} + \frac{1}{r}\frac{\partial}{\partial r} + \frac{\partial^2}{\partial z^2}$ ,  $u_r$  and  $u_z$  are the radial and vertical displacement components, respectively,  $\varepsilon_v = \frac{\partial u_r}{\partial r} + \frac{u_r}{r} + \frac{\partial u_z}{\partial z}$  is the volumetric strain.

It is assumed that the solid constituent is incompressible but the fluid constituent is compressible. Combined with Darcy's law, the pore fluid mass conservation equation is given by

$$\frac{k}{\gamma_w}\nabla^2 p = \frac{\partial}{\partial t}\left(\varepsilon_v + \frac{p}{M}\right) \quad (6)$$

where  $k$  is the coefficient of permeability of the soil,  $\gamma_w$  is the unit weight of water,  $M$  is the bulk modulus (adjusted for porosity) of the pore fluid, and  $t$  is a time variable.

### 2.2. Boundary conditions

Fig. 1 illustrates a groundwater drawdown in an aquifer. The aquifer is assumed to be rough and rigid, so there is no vertical or radial displacement at the bottom boundary of the soil layer. The upper boundary of the soil layer is permeable and stress-free. The boundary conditions can be expressed as

$$z = 0 : \sigma_{zz} = 0, \sigma_{rz} = 0, p = 0. \quad (7)$$

$$z = H : u_z = 0, u_r = 0, p = -\gamma_w h(r, t). \quad (8)$$

## 3. Solutions

### 3.1. General solutions

Laplace and Hankel transforms were employed to obtain the solutions. The  $n$ th-order Laplace–Hankel transform of  $f(r, z, t)$  and its inversion may be found for example in Sneddon [21]:

$$\hat{f}(\xi, z, s) = \int_0^\infty \int_0^\infty f(r, z, t) r J_n(\xi r) e^{-st} dr dt \quad (9)$$

$$f(r, z, t) = \frac{1}{2\pi i} \int_{\gamma-i\infty}^{\gamma+i\infty} \int_0^\infty \hat{f}(\xi, z, s) \xi J_n(\xi r) e^{st} d\xi ds \quad (10)$$

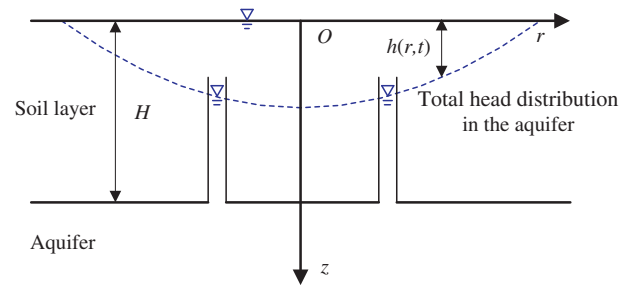


Fig. 1. Consolidation of a soil layer induced by groundwater drawdown in an aquifer.

Download English Version:

<https://daneshyari.com/en/article/254792>

Download Persian Version:

<https://daneshyari.com/article/254792>

[Daneshyari.com](https://daneshyari.com)