



# Quantification of the gas mass emitted during single explosions on Stromboli with the SO<sub>2</sub> imaging camera

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## ABSTRACT

We performed measurements using an SO<sub>2</sub> imaging camera of the SO<sub>2</sub> gas mass emitted during five discrete explosive events on Stromboli volcano on 3 October 2006. The SO<sub>2</sub> gas mass released during discrete explosions was 15–40 kg per explosion, producing 3–8% of the total daily SO<sub>2</sub> gas emission, demonstrating that in terms of gas flux Strombolian explosions are a second-order phenomenon compared with quiescent degassing. Using the typical gas composition measured with OP-FTIR allows us to determine the total gas mass released during an explosion as 360–960 kg with a volume of 1500–4100 m<sup>3</sup> at 1 bar. At the probable source pressure of gas slug formation of 75 MPa this gas amount would occupy a volume equivalent to a sphere with a radius of 0.8–1 m, comparable with estimates of Stromboli's conduit geometry.

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## 1. Introduction

Stromboli, the most easterly volcanic island in the Aeolian island archipelago, lies ~60 km north of Sicily, Italy, (see Fig. 1) and is one of the most active volcanoes in the world. The activity of the volcano is characterized by persistent degassing and intermittent mild eruptions ejecting incandescent bombs and gas, so called Strombolian eruptions (Barberi et al., 1993). The dynamics of Strombolian eruptions has been studied extensively by seismicity, geochemistry, infrasound and petrology by various researchers (e.g. Ripepe et al., 2002; Bertagnini et al., 2003; Chouet et al., 2003; Burton et al., 2007a). Recent advances in hand-held forward looking infrared radiometer (FLIR) cameras has allowed detailed observation of eruptive plume dynamics of Stromboli (Patrick et al., 2007; Patrick, 2007; Harris and Ripepe, 2007).

Stromboli is an ideal target to understand the relationship between eruptive activity, chemical composition and flux of volcanic gas (Allard et al., 1994). Although its persistent explosive activity hinders direct sampling of high temperature volcanic gas directly from the crater vents, the challenge of monitoring chemical composition has been recently surmounted by using new techniques such as FTIR (Mori and Notsu, 1997; Allard et al., 2005; Burton et al., 2007a) and MULTIGAS (Shinohara, 2005; Aiuppa et al., 2009). Burton et al. (2007a) successfully showed that volcanic gas at Stromboli for quiescent degassing and explosions has different chemical compositions and that swiftly ascending gas slugs are responsible for the

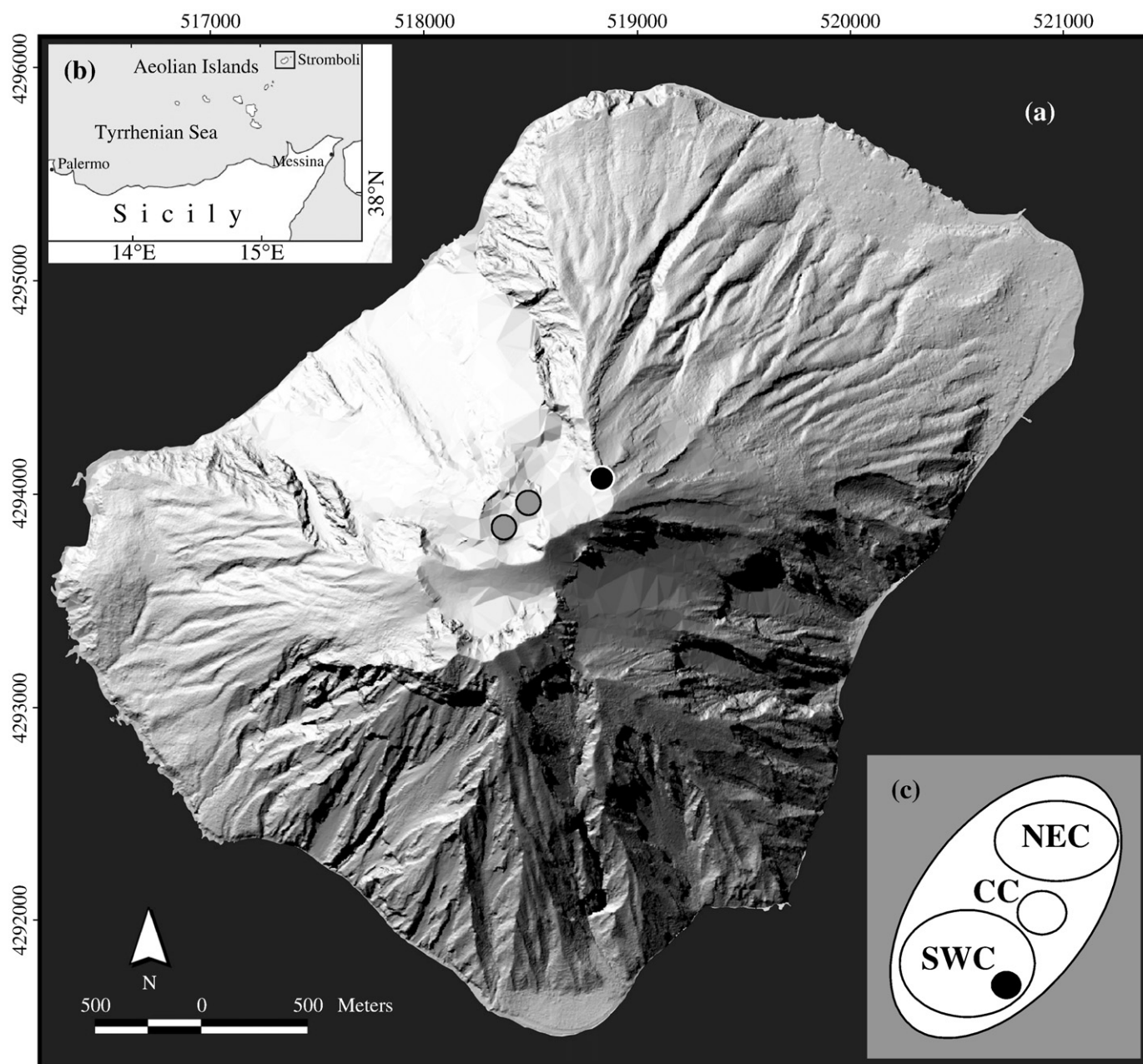
explosive activity, rising from the volcano-crust interface at about 3 km depth.

Volcanic gas flux carries important information on both the amount of magma degassing and the magma plumbing system. Although the major volcanic gas component is water vapour, SO<sub>2</sub> has been the target of gas flux measurements (Stoiber et al., 1983; Galle et al., 2002; McGonigle et al., 2002) because of easy detection by remote ultraviolet spectroscopy and low background concentration. On Stromboli, fluxes have been found to range between 100 t/day and 1400 t/day with mean of several hundred tons per day of SO<sub>2</sub> (Allard et al., 1994; Allard et al., 2008). Recent installation of scanning ultraviolet spectrometers on the island has enabled automatic measurement of SO<sub>2</sub> flux at high temporal resolution (Burton et al., 2009), and the detection of elevated fluxes prior to paroxysmic explosions.

While Strombolian eruptions occur with a frequency of 8–17 events/h at Stromboli (Harris and Ripepe, 2007), there are relatively few measurements of the gas flux released during the explosions. Allard et al. (1994) used air-borne COSPEC to detect a 40–400% increase in SO<sub>2</sub> flux during explosions relative to the quiescent flux. Several geophysical approaches have attempted to constrain the volume of gas involved in explosive activity; photoballistic measurements of jets (600–1200 m<sup>3</sup> gas/explosion, Chouet et al., 1974), acoustic measurements (1–100 m<sup>3</sup> gas/explosion, Vergnolle and Brandeis, 1996) and explosion quake modelling (~200 m<sup>3</sup> gas/explosion at source pressure, Chouet et al., 2003). To date there have been no direct volcanic gas measurements constraining the SO<sub>2</sub> amount emitted by a single Strombolian eruption. This is an important parameter, as it constrains the relative amount of gas lost quiescently compared with explosions, and highlights quantitatively the relative

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**Fig. 1.** (a) Map of Stromboli with the Rocchetta observation site (black circle) and summit craters (grey circles) indicated. Approximate measurement distance is 650 m. (b) Inset indicating location of Stromboli with respect to Sicily. (c) Schematic map of the crater area, with Northeast crater (NEC), central crater (CC), Southwest crater (SWC) and hornito (black circle) indicated.

importance of each process in terms of magma supply and degassing. In this paper, we present the first measurements of  $\text{SO}_2$  amounts emitted by single Strombolian events using the recently developed  $\text{SO}_2$  imaging camera (Mori and Burton, 2006; Bluth et al., 2007).

## 2. $\text{SO}_2$ imaging camera measurements at the summit of Stromboli

The  $\text{SO}_2$  imaging camera (Mori and Burton, 2006) consists of a UV-sensitive CCD camera (Apogee Alta U-260), C-mount quartz lens (Pentax B2528-UV) and UV bandpass filters (Asahi Bunko Co.). The camera sensor is a  $512 \times 512$  pixel 16-bit CCD. In the original configuration of the  $\text{SO}_2$  camera, two UV bandpass filters centred at 310 nm and 330 nm with FWHM of  $\sim 10$  nm were used for image processing (Mori and Burton, 2006). The 310 nm filter is sensitive to both  $\text{SO}_2$  absorption and aerosol attenuation, whereas the 330 nm filter is only sensitive to aerosol attenuation. Obtaining images

alternately using the two filters allowed the aerosol attenuation to be removed, revealing the uncontaminated  $\text{SO}_2$  signal.

On the afternoon of October 3, 2006, we took the  $\text{SO}_2$  camera to the summit area of Stromboli and placed it at the Rocchetta site (Fig. 1a), about 450 m from and at the same altitude as the crater terrace ( $\sim 750$  m), where the ascending plume emitted from the summit craters could be observed against the background sky. The NE crater was directly visible and Central and SW craters were hidden behind the NE crater. The eruption plumes of the Central and SW craters were only observed when they rose above the rim. On the South-eastern edge of the SW crater, there was a hornito several tens of meters high. The top of this hornito was directly visible from the Rocchetta site above the crater terrace rim (Fig. 2).

At the time of observation, all the craters including the hornito were persistently degassing. The NE crater was the most vigorously degassing vent with puffing activity similar to the puffing conditions

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