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# Rapid measurements of heterogeneity in sandstones using low-field nuclear magnetic resonance

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#### ABSTRACT

Sandstone rocks can contain microscopic variations in composition that complicate interpretation of nuclear magnetic resonance (NMR) relaxation time measurements. In this work, methods for assessing the degree of sample heterogeneity are demonstrated in three sandstones. A two-dimensional  $T_1 - \Delta \chi_{app}$  correlation (where  $\Delta \chi_{app}$  is the apparent solid/liquid magnetic susceptibility contrast) reveals the microscopic heterogeneity in composition, whilst a spatially resolved  $T_1$  profile reveals the macroscopic structural heterogeneity. To perform these measurements efficiently, a rapid measure of longitudinal  $T_1$  relaxation time has been implemented on a low-field NMR spectrometer with a magnetic field strength  $B_0 = 0.3$  T. The "double-shot"  $T_1$  pulse sequence is appropriate for analysis of porous materials in general. Example relaxation time of the double-shot  $T_1$  sequence is equivalent to the single-shot Carr-Purcell Meiboom–Gill (CPMG) sequence used routinely in petrophysics to measure transverse  $T_2$  relaxation. Rapid  $T_1$  measurements enable practical studies of core plugs at magnetic field strengths previously considered inappropriate, as  $T_1$  is independent of molecular diffusion through pore-scale (internal) magnetic field gradients.

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## 1. Introduction

Nuclear magnetic resonance (NMR) relaxation times are sensitive to the liquid/solid interface in porous materials. Transverse  $T_2$  relaxation time measurements are used extensively in petrophysical rock core analysis [1,2]. The Carr–Purcell Meiboom–Gill (CPMG) pulse sequence [3,4] provides a rapid, single-shot method for measuring the exponential decay of transverse magnetization. In petrophysics, the CPMG sequence is most usefully applied at low magnetic field strengths,  $B_0 \ll 0.1$  T, to minimize the effects of magnetic susceptibility contrast in these heterogeneous samples. A  $T_2$  distribution has become the archetypal NMR well logging measurement, where analysis of the data provides estimates of porosity, permeability, and saturation state of the rock formation near the well bore [5]. When the pore structure contains only a single liquid, a relaxation time distribution is considered equivalent to a pore size distribution [6], assuming a single surface relaxivity parameter controls the scaling between relaxation time and pore size. Historically, sandstones have been studied by NMR because of the simplicity of their granular structure, although the mineralogy of sandstones will often be complicated. When the sandstone contains spatially localized mineral deposits (e.g.,

ties will exist, which complicates the interpretation of relaxation time distributions [7]. NMR measurements are sensitive to variations in local sample composition through (a) surface relaxivity and/or (b) magnetic sus-

ceptibility contrast  $\Delta \chi$  between the solid pore matrix and the sat-

feldspar, clays, or metal oxides), a distribution of surface relaxivi-

urating liquid. The  $T_2$  measurement is influenced by both surface relaxation and diffusion through pore-scale "internal" magnetic field gradients caused by the susceptibility contrast [8]. Diffusive signal attenuation results in the observation of an effective relaxation time,  $T_{2,eff}$ . The magnitude of the internal gradients is determined by  $\Delta \chi B_0$ , so minimizing the magnetic field strength reduces the influence of diffusion on the  $T_2$  measurement. Laboratory-scale core analysis is routinely performed at  $B_0 = 50$  mT, corresponding to  $v_0 = 2$  MHz for <sup>1</sup>H, (a) to minimize internal gradient effects and (b) to ensure robust calibration of well logs acquired at a similar Larmor frequency. The sensitivity of NMR experiments to microscopic variations in composition increases with magnetic field strength as both surface

composition increases with magnetic field strength as both surface relaxivity [9] and internal gradients [8] are field dependent. In the present work, a  $B_0 = 0.3$  T magnet is used to enhance sensitivity to structure in the NMR experiment. This field strength was previously considered inappropriate for core analysis of sandstones because of the effects of internal gradients [10].







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Notwithstanding, quantitative information is obtained through measurements of the longitudinal  $T_1$  relaxation time.  $T_1$  was the original metric of choice in NMR petrophysics [6,11] due to hardware constraints in well logging tools and the robustness of the  $T_1$  measurement [10,12]. In well logging,  $T_1$  was later rejected in favor of  $T_2$ : the single-shot CPMG measurement has sufficient time resolution to be useful on moving tools [13]. Laboratory-scale core analysis followed this trend, and now low-field measurements of  $T_2$  are considered the industry standard.

In the laboratory context, it is usual to acquire  $T_1$  data using the conventional inversion or saturation recovery protocols [14]. However, several rapid  $T_1$  measurement protocols are available, as reviewed in [15] with recent method developments given in [16–22]. The double-shot  $T_1$  measurement has an acquisition time comparable to that of the CPMG sequence [23]. In the double-shot  $T_1$  experiment, the relaxation time is encoded as an exponential decay following two radio frequency (rf) excitation pulses. In the present work, the double-shot  $T_1$  experiment (see Section 2) is validated for use on a low-field magnet. Quantitative  $T_1$  distributions are presented for doped water phantoms (see Section 3) and for sandstone core plugs (see Section 4).

To determine microscopic heterogeneity in sample composition (mineralogy), a  $T_1\text{-}\Delta\chi_{app}$  correlation is introduced.  $\Delta\chi_{app}$  is an apparent solid/liquid magnetic susceptibility contrast, obtained from a measurement of  $T_2^*$ , the transverse relaxation rate determined by local field inhomogeneities [24,25]. An apparent susceptibility contrast is obtained because the measured  $T_2^*$  is influenced by details of the local pore geometry as well as the local magnetic field [25]. Localized variations in composition (occurring over length-scales much greater than the diffusion path-length of water) should result in observable variations in  $\Delta \chi$ . Furthermore, these same localized changes in composition will provide variations in surface relaxivity.  $\Delta \chi_{app}$  (through  $T_2^*$ ) and  $T_1$  offer independent probes of these two composition-dependent parameters. A  $T_1 - \Delta \chi_{app}$  correlation therefore offers a straight-forward visual representation of heterogeneity: a strong inverse correlation between the two dimensions is indicative of compositional variation. Elsewhere,  $T_1$  has been correlated against effective internal gradient strength  $g_{\text{eff}}$  [12], whilst  $T_{2,\text{eff}}$  has been correlated against  $\Delta \chi$  [26]. However, in the former experiment,  $g_{\rm eff}$  is difficult to determine in general for sandstones [27], and in the latter experiment, both dimensions are dependent on the internal gradient distribution. The  $T_1$ - $\Delta \gamma$  correlation provides direct information on the microscopic compositional heterogeneity as it unambiguously correlates magnetic susceptibility contrast with surface relaxation.

Macroscopic structural heterogeneity is explored by incorporating imaging gradients into the double-shot  $T_1$  sequence to enable rapid acquisitions of  $T_1$  profiles. Spatially resolved  $T_1$  measurements have been used previously to study drilling mud filter cakes [28] and the invasion of clays and drilling mud into core plugs [28–30] where the presence of paramagnetic species (e.g., iron in barite used as a densifying agent in mud) makes conventional  $T_2$  analysis inappropriate [1]. Rapid methods of mapping  $T_1$  will enable time-resolved studies of such invasion processes and also permit monitoring of oil recovery [31] at magnetic field strengths previously considered inappropriate for petrophysical core analysis.

### 2. Experimental methods

### 2.1. Pulse sequences

All NMR measurements were performed on an Oxford Instruments  $B_0 = 0.3$  T vertical bore, bench-top permanent magnet operating at  $v_0 = 12.9$  MHz for <sup>1</sup>H. Imaging gradients up to  $g_{x,y,z}^{(max)} = 25$  G cm<sup>-1</sup> were provided by three-axis slab-format gradient coils. The gradient coils were driven by high-performance ERA audio amplifiers. A field homogeneity of  $\Delta v_0 = 5$  ppm was achieved over the entire sample volume (cylindrical rf probe of dimensions length  $\times$  diameter: 6.0 cm  $\times$  5.1 cm) by high-order shims. The experiments were controlled by an Oxford Instruments Maran DRX-HF spectrometer using an interface written in the Matlab programming environment.

The original double-shot  $T_1$  pulse sequence, Fig. 1(a) [23], was based in part on the  $T_1$  by multiple read-out pulses (TOMROPs) sequence [32]. However, where the TOMROPs sequence generates a relaxation curve defined by a complicated function of  $T_1$ , the total magnetization  $M_0$ , and the driven equilibrium magnetization  $M_{\infty}^{eq}$ (following an "infinite" number of small tip angle rf pulses), the double-shot sequence generates a simple exponential decay that depends only on  $T_1$  and  $M_0$ . In [23], a 90°–180°–90° preconditioning sequence was used to store the magnetization alternately along  $\pm z$  so that after two scans, the signal amplitude was simply  $b(t) \propto M_0 \exp\{-t/T_1\}$ , where *t* is the total relaxation time.

Here, a modified version of the double-shot  $T_1$  sequence is introduced, as shown in Fig. 1(b). The first portion of the pulse sequence consists of a series of small tip angle  $\alpha$  pulses (such that sin  $\alpha \approx \alpha$ ), each pulse being separated in time by  $\tau_1$ . A complete free induction decay (FID) [33] is acquired immediately after each  $\alpha$  pulse. This portion of the pulse sequence measures  $M_{\infty}^{eq}$  obtained during the train of *n* pulses. A comb of 90° pulses is then applied to fully saturate the spin ensemble in the *x*-*y* plane. A second train of  $\alpha$  pulses is used to observe the recovery of the magnetization from zero to  $M_{\infty}^{eq}$  on the *z*-axis. In a post-processing stage, the driven equilibrium signal (acquired in the first portion) is subtracted from the recovery curve (acquired in the second portion) to provide an exponential relaxation decay. The normalized signal amplitude is then given by

$$\frac{b(t)}{b(0)} = \exp\left\{-\frac{t}{T_1}\right\} (\cos\alpha)^{n-1} \sin\alpha, \tag{1}$$

with

$$t = (n-1)\tau_1 + t_{\text{spoil}},\tag{2}$$

where  $t_{\text{spoil}}$  is the 90°- $\alpha$  delay required to apply a homospoil gradient pulse. To prevent the refocusing of unwanted magnetization, homospoil gradient pulses of variable amplitude are included between each  $\alpha$  pulse. To guarantee quantitative signal amplitudes, it is only necessary to provide a recycle delay of  $t_{\text{RD}} \ge \tau_1$  between each acquisition if the magnetization has recovered to  $M_{\text{RD}}^{\text{eq}}$  at the end of the acquisition. Otherwise, the condition  $t_{\text{RD}} \gg \tau_1$  is required. The modified double-shot  $T_1$  sequence has a slightly reduced acquisition time ( $t_{\text{RD}}$  is included only once every two scans) compared to the original implementation ( $t_{\text{RD}}$  is included between every scan). For practical implementation on low-field hardware,



**Fig. 1.** Double-shot  $T_1$  pulse sequence schematic. The original pulse sequence (a) was implemented at high field, and the modified pulse sequence (b) is used here at low field. The thin and thick vertical lines indicate 90° and 180° rf pulses, respectively, with the small tip angle ( $\alpha$ ) pulses indicated by vertical lines of reduced amplitude. The gray trapezoids indicate homospoil gradient pulses. The time between successive  $\alpha$  pulses is  $\tau_1$ , and the 90°- $\alpha$  delay is  $t_{spoil}$ . The half-echo time in (a) is  $\tau_2$ .

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