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# DNA-based stable isotope probing identifies formate-metabolizing methanogenic archaea in paddy soil



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#### ABSTRACT

Paddy methane (CH<sub>4</sub>) production is biologically dominated by methanogenic archaea that metabolize a variety of organic and/or inorganic carbon sources. Though formate is easily dissimilated into  $H_2/CO_2$ , formate-metabolizing methanogenic archaea are distinct from  $CO_2$ -utilizing methanogen taxa. The identity of formate-metabolizing methanogenic archaea in paddy soil remains elusive. In this investigation, molecular approaches based on stable isotope probing (SIP) technique were conducted to identify the formate-metabolizing methanogenic archaea in paddy soil.  $CH_4$  emission monitor, real-time quantitative PCR (qPCR) and Denaturing Gradient Gel Electrophoresis (DGGE) analyses consistently indicated that some methanogenic archaea metabolized  $^{13}$ C-labeled formate in microcosm and accounted for a large portion of formate-metabolizing archaea in anoxic paddy soil. Phylogenetic identification further found that this guild was affiliated to *Methanobacteriaceae*. Taken together *Methanobacteriaceae* could be the dominant formate-metabolizing methanogenic archaea and play an important role in the  $CH_4$  production in paddy soil. These findings would extend the extant knowledge on paddy methanogenic archaea and microbial-driven paddy  $CH_4$  emission.

#### 1. Introduction

The global warming potential of methane (CH<sub>4</sub>) is 25 times greater than that of carbon dioxide (CO<sub>2</sub>) on 100-year horizon. IPCC (2007) reported that among global anthropogenic greenhouse gas emissions in 2004, CH<sub>4</sub> accounts for 14.3% in terms of CO<sub>2</sub> equivalent. Rice is the world's most important agronomic plant, with nearly 150 million ha under cultivation globally (Roger, 1996). In about 75% of this land rice grows under flooded conditions. Thus, paddy fields are one of largest anthropogenic sources of global CH<sub>4</sub> emissions (IPCC, 2007).

Biologically, paddy CH<sub>4</sub> production is dominated by methanogenic archaea. Methanogenesis is the final degradation process of organic matter in paddy soil and dependent on intermediary substrates derived from their degradation. Organic matter is first anoxically degraded to small molecules, such as acetate, formate, CO<sub>2</sub> and H<sub>2</sub>, by diverse bacteria. Methanogenic archaea metabolize some of these molecules and further convert into CH<sub>4</sub> (Watanabe et al., 2007). Acetate contributes 79–83% to CH<sub>4</sub> production as a carbon substrate for methanogens (Chin and Conrad, 1995). In paddy soil, the CH<sub>4</sub> produced from acetate was between 51- 67% (Chidthaisong et al., 1999), while other

substrates, such as  $\rm H_2/CO_2$  (or formate) contribute 17 - 31% to CH<sub>4</sub> production (Rothfuss and Conrad, 1992). Therefore, the identification of specific substrate-metabolizing methanogenic archaea is of great significance toward the knowledge of both the microbial-driven paddy carbon cycle and ecological functions of paddy methanogenic archaeal guild.

The CO<sub>2</sub>- and acetate-utilizing methanogens in anoxic paddy soil have been previously documented. For example, Liesack et al. (2000) summarized that the predominant acetate-utilizing methanogens in paddy soil belong to *Methanosarcinaceae* and *Methanosaetaceae*. Members of the *Methanobacteriaceae* are the predominant group of hydrogenotrophic methanogens in rice paddy soils (Tonouchi, 2002). However, the information regarding identity of formate-assimilating methanogenic archaea in paddy soil remains limited. Formate is an important organic acid and a significant driver of methanogenesis in paddy soil (Penning and Conrad, 2006). The concentration of formate can exceed 150 µM in anoxic paddy soil (Rothfuss and Conrad, 1992). In spite of the easy dissimilation of formate to H<sub>2</sub> and CO<sub>2</sub>, we propose that formate can be directly metabolized by methanogenic archaea to produce CH<sub>4</sub> in paddy soils, and the formate- and CO<sub>2</sub>-utilizing

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Abbreviations: DGGE, denaturant gradient gel electrophoresis; DNA-SIP, DNA-based stable isotope probing; FACE, free-air CO<sub>2</sub> Enrichment; qPCR, real-time quantitative PCR; SOC, soil organic C

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methanogen taxa are presumably distinct. Indeed, Hunger et al. (2011) have identified different functional guilds in the actively formate-metabolizing and  $\rm H_2/CO_2$ -utilizing methanogens in fen soil. However, there is no unambiguous evidence to identify formate-assimilating methanogens in flooded paddy soil. DNA-based stable isotope probing (DNA-SIP) is a helpful technique to link the microbial metabolic function and their taxonomic identity in complex environment (Radajewski et al., 2003). The technique has been used to identify dominant formate-metabolizing bacteria in paddy soil, phylogenetically related to the bacteria in the classes of Clostridia and  $\alpha$ -Proteobacteria (Feng et al., 2012), and to distinguish the difference in actively formate-assimilating and  $\rm H_2/CO_2$ -utilizing methanogens in fen soil (Hunger et al., 2011). Therefore, in this investigation a microcosm-based DNA-SIP experiment was employed to identify formate-metabolizing methanogens in a flooded paddy soil.

#### 2. Materials and methods

#### 2.1. Soil sampling

Soil samples were taken from a rice-wheat rotation paddy field in Jiangsu Province, China (31°35′N, 120°30′E). The soil is classified as stagnic anthrosols. The soil properties are shown in Table 1. The station sits in the subtropical climatic zone with a mean annual precipitation of 900–1000 mm, the mean air temperature between June and August of 30 °C, an average daily integral radiation of 12.3 MJ/m², a total annual sunshine time of more than 2000 h and a frost-free period of more than 230 days. Soil samples at depth of 0–5 cm were collected from five points along S curve. Then the samples were kept in a cooler and shipped to the lab as quickly as possible. In the lab, the soil samples were pooled, passed through a 2 mm sieve and immediately stored at 4 °C after plant materials, roots, and stones were removed.

#### 2.2. Anoxic microcosms

The DNA-SIP microcosm contained 5 g bulk soil and 0.5 mmol of  $^{13}\text{C}$ -labeled formate (99 atom at%  $^{13}\text{C}$ ) purchased from Sigma-Aldrich (Kalyuzhnaya et al., 2008). As comparisons, the microcosms of  $^{12}\text{C}$ -formate addition and the control without formate addition were established. Each treatment had three replicates. All the incubations were performed in sterile 120 ml serum bottle. The serum bottles were closed with rubber stoppers and crimp seals and flushed with sterile  $N_2$  (100%) to make microcosms anaerobic. A 24-day anoxic incubation of microcosms was performed at 60% soil maximum water holding capacity, 30 °C and darkness. Gas samples (10 ml) were taken with a gastight syringe from the headspace of the soil microcosms every 3 days. The same volume of  $N_2$  was then injected into the bottle to keep the equitant pressure. Totally, 24 gas samples were collected for each treatment. Production of CH4 was monitored by a Varian 3380 with FID.

### 2.3. DNA extraction and SIP gradient fractionation

On the day after final gas sampling, soil samples from each microcosm were collected, mixed and sieved ( <2 mm). Samples were kept at  $-20\,^{\circ}\text{C}$  for molecular analysis. A half gram of moist soil from each sample was used for DNA extraction using FastDNA $^{\circ}$  SPIN Kit for Soil (MP Biomedicals, Santa Ana, CA) according to the manufacturer's instructions. The extracted DNA was dissolved in 50  $\mu$ l TE buffer, quantified by spectrophotometer and stored at  $-20\,^{\circ}\text{C}$  until further use.

DNA-SIP was performed by published protocol (Hunger et al., 2011). Briefly, the gradient fractionation of total DNA extract (3.0  $\mu$ g) from each SIP microcosm was performed with an initial CsCl buoyant density of 1.720 kg l $^{-1}$  subjected to centrifugation at 177,000g for 44 h at 20 °C. The density gradient was divided into 340- $\mu$ l fractions and the buoyant density of each fraction was determined by the refractive

index. Fifteen fractions were generated covering buoyant densities from  $1.696\ kg\ l^{-1}$  to  $1.743\ kg\ l^{-1},$  and nucleic acids were separated from cesium chloride by PEG 6000 precipitation and the resulting pellets were dissolved in 30  $\mu l$  of TE buffer.

#### 2.4. Real-time quantitative PCR

The abundances of methanogenic archaeal (primer set 1106F/ 1378R) and total archaeal (primer set A364aF/A934b) 16S rRNA genes along density gradients were quantified by quantitative PCR (qPCR) following the protocols of Watanabe et al. (2007) and Kemnitz et al. (2005) respectively. Standard curves were obtained using 10-fold serial dilutions of the Escherichia coli-derived vector plasmid pMD18-T (Ta-KaRa) containing a cloned corresponding target gene, using 10<sup>2</sup> to 10<sup>8</sup> gene copies µl<sup>-1</sup>. The reactions were performed in C1000™ Thermal Cycler equipped with CFX96™ Real-Time system (Bio-Rad, USA). The 25-μl reaction mixture contained 12.5 μl of SYBR® Premix Ex Taq™ (TaKaRa), primer set (0.5  $\mu$ M each), 200 ng BSA  $\mu$ l<sup>-1</sup>, 1.0  $\mu$ l template containing approximately 2-9 ng DNA. Negative control was always run with water as the template instead of soil DNA extract. The qPCR program used for methanogenic archaea or total archaea was: 94 °C for 5 min, followed by 35 cycles of 94 °C for 30 s, 55 °C or 66 °C for 30 s and 72 °C for 90 s or 50 s, and extension and signal reading. The specificity of the amplification products was confirmed by melting curve analysis, and the expected sizes of the amplified fragments were checked in a 1.5% agarose gel. Real-time qPCR was performed in triplicate and amplification efficiencies of 97.4-104% were obtained with  $R^2$  values of 0.966–0.977.

#### 2.5. Denaturant gradient gel electrophoresis (DGGE) analysis

Methanogenic archaeal 16S rRNA gene fragments in 280-bp length were amplified using the primer set 1106F-GC (CGCCGCGCGCGCGCG GCGGGCGGG CGGGG GCACGGGGGGTTWAGT CAG GCAACGAGC) and 1378R (CCCATGGTCCAGC GCCAGAA) (Watanabe et al., 2007)along buoyant density gradientsfor all treatments. Approximately 150-250 ng PCR amplicons from each sample were loaded onto an 8% (w/v) acrylamide-bisacrylamide gel with 45%-75% denaturant gradient. DGGE was run in 1  $\times$  TAE buffer for 10 h at 60  $^{\circ}\text{C}$  and 100 V with a Dcode Universal Mutation Detection System (Bio-Rad, Hercules, Calif.). After electrophoresis, the DGGE gel was stained for 20 min with SYBR Green I nucleic acid staining solution with 1:10000 (v/v) (Invitrogen, Oregon, USA) and photographed. DGGE fingerprinting profiles of methanogenic archaeal 16S rRNA genes were digitalized by using a Gel Doc™ EQ imager (Bio-Rad, USA) combined with Quantity One 4.4.0 (Bio-Rad, Hercules, CA, USA). The dominant and representative bands in DGGE gels were excised, left overnight in 25 µl Milli-Q water, re-amplified and run again on the DGGE system to ensure purity and correct mobility of the excised DGGE bands.

#### 2.6. Cloning, sequencing and phylogenetic analysis

After confirmation, the excised DGGE bands were re-amplified with the primer set of methanogenic archaea without GC clamp, followed by the purification using the QIAquick PCR Purification Kit (QIAGEN). The purified PCR products were inserted into a pMD18-T vector (TaKaRa) in accordance with the manufacturer's instructions and further introduced into *Escherichia coli* DH5 $\alpha$  competent cell. Six random clones were sequenced by Invitrogen Sequencing Department in Shanghai, China.

One representative clone sequence of each band with high quality after sequence comparisons using DNASTAR software package was chosen for phylogenetic analysis. The representative sequences of DGGE bands were compared with sequences in BLAST to obtain the three nearest phylogenetic neighbors. Then a phylogenetic tree was built by the neighbor-joining method using the software package of MEGA 4.0 version (Molecular Evolutionary Genetics Analysis) (Tamura et al., 2007). The GenBank accession numbers for methanogenic

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