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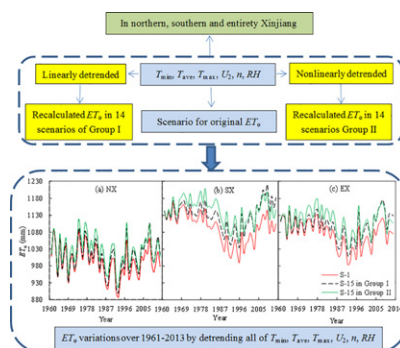
# Influences of removing linear and nonlinear trends from climatic variables on temporal variations of annual reference crop evapotranspiration in Xinjiang, China

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## HIGHLIGHTS

- Removing climatic variable trends increased reference crop evapotranspiration ( $ET_0$ ).
- $ET_0$  changes were more convincing when climatic variables were nonlinearly detrended.
- $ET_0$  obtained from the detrended CVs caused different variations of drought severity.

## GRAPHICAL ABSTRACT



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## ABSTRACT

Reference crop evapotranspiration ( $ET_0$ ) is a key parameter in field irrigation scheduling, drought assessment and climate change research.  $ET_0$  uses key prescribed (or fixed or reference) land surface parameters for crops. The linear and nonlinear trends in different climatic variables (CVs) affect  $ET_0$  change. This research aims to reveal how  $ET_0$  responds after the related CVs were linearly and nonlinearly detrended over 1961–2013 in Xinjiang, China. The  $ET_0$ -related CVs included minimum ( $T_{\min}$ ), average ( $T_{\text{ave}}$ ), and maximum air temperatures ( $T_{\max}$ ), wind speed at 2 m ( $U_2$ ), relative humidity (RH) and sunshine hour ( $n$ ).  $ET_0$  was calculated using the Penman-Monteith equation. A total of 29  $ET_0$  scenarios, including the original scenario, 14 scenarios in Group I ( $ET_0$  was recalculated after removing linear trends from single or more CVs) and 14 scenarios in Group II ( $ET_0$  was recalculated after removing nonlinear trends from the CVs), were generated. The influence of  $U_2$  was stronger than influences of the other CVs on  $ET_0$  for both Groups I and II either in northern, southern or the entirety of Xinjiang. The weak influences of increased  $T_{\min}$ ,  $T_{\text{ave}}$  and  $T_{\max}$  on increasing  $ET_0$  were masked by the strong effects of decreased  $U_2$  and  $n$  and increased RH on decreasing  $ET_0$ . The effects of the trends in CVs, especially  $U_2$ , on changing  $ET_0$  were clearly shown. Without the general decreases of  $U_2$ ,  $ET_0$  would have increased in the past 53 years.

**Abbreviations:** CV, climatic variable;  $ET_0$ , reference crop evapotranspiration;  $T$ , air temperature;  $T_{\min}$ , minimum  $T$ ;  $T_{\text{ave}}$ , mean  $T$ ;  $T_{\max}$ , maximum  $T$ ;  $U$ , wind speed;  $U_2$ ,  $U$  at 2 m; RH, relative humidity;  $n$ , sunshine hour;  $P$ , precipitation; NX, northern Xinjiang; SX, southern Xinjiang; EX, entire Xinjiang; EEMD, ensemble empirical mode decomposition; MK, Mann-Kendall; MMK, modified MK; Z, MK statistic;  $Z_{\text{m}}$ , MMK statistic; SQMK, sequential Mann-Kendall; S, scenario;  $b$ , Sen's slope; RS, regression slope;  $Y_{\text{AC}}$ , year with an abrupt change;  $P_{\text{sl}}$ , significant level;  $G$ , soil heat flux;  $e_a$ , actual vapor pressure;  $e_s$ , saturation vapor pressure;  $\Delta$ , slope of vapor pressure curve;  $\gamma$ , psychrometric constant;  $R_n$ , net radiation; SD, standard error, RMSE-root mean square error.

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Due to the non-monotone variations of the CVs and  $ET_o$ , the results of nonlinearly detrending CVs on changing  $ET_o$  in Group II should be more plausible than the results of linearly detrending CVs in Group I. The decreasing  $ET_o$  led to a general relief in drought, which was indicated by the recalculated aridity index. Therefore, there would be a slightly lower risk of water utilization in Xinjiang, China.

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## 1. Introduction

The global average land surface temperature increased by 0.85 °C in the period from 1880 to 2012 (IPCC, 2013). Faced with the challenges of global warming, it is crucial to know how the water availability and the agricultural water demands change, as a function of reference crop evapotranspiration ( $ET_o$ ) variation (Espadafor et al., 2011).  $ET_o$  is widely used for water scheduling of irrigation areas and uses key prescribed (or fixed or reference) land surface parameters for crops. Accurate estimation of  $ET_o$  is needed for field irrigation scheduling (Allen et al., 1998), drought assessment (Budyko, 1974) and climate change research (Tabari and Hosseinzadeh Talaei, 2014). The available models for estimating  $ET_o$  include temperature-based approaches (Blaney and Criddle, 1952; Hargreaves and Samani, 1985; Thornthwaite, 1948), radiation-based approaches (Slatyer and McIlroy, 1961) and combination approaches (Allen et al., 1998; Priestley and Taylor, 1972). Following an extensive analysis of different  $ET_o$  methods in many world-wide locations, the Penman-Monteith (PM) approach was unanimously accepted as a standardized approach endorsed by the Food and Agricultural Organization (Allen et al., 1998; McVicar et al., 2007).

In the context of global warming, the trend of  $ET_o$  was found to have increased or decreased in different regions of the world. Increasing trends of  $ET_o$  were observed in Iran (Tabari et al., 2011, 2012) and southern Spain (Espadafor et al., 2011). However, a decreasing  $ET_o$  trend was reported for India (Chattopadhyay and Hulme, 1997).  $ET_o$  at many places in China have decreased (see detail in Table 6 from McVicar et al., 2007). Trends in  $ET_o$  have decreased in northwest and southeast regions (Thomas, 2000), the Yangtze River Basin (Xu et al., 2006), part of the Yellow River Basin (Liu and Yang, 2010), Tibet Plateau (Liu et al., 2011), and the Haihe River Basin (Tang et al., 2011). The increase or decrease of  $ET_o$  were mainly caused by the increase or decrease of climatic variables (CVs), which are necessary for estimating  $ET_o$ . These variables include air temperature ( $T$ ), wind speed ( $U$ ), relative humidity ( $RH$ ), and radiation or sunshine hour ( $n$ ) in the absence of radiation. The influences of CVs on  $ET_o$ , which also potentially reflect the effects of climatic change on  $ET_o$ , have been investigated regionally or globally (McVicar et al., 2012; Xu et al., 2006).

There are different ways for assessing the influences of CVs on  $ET_o$ . One way is to conduct a sensitivity analysis. This is done by determining the corresponding changes of  $ET_o$  after adding different percentages of the CVs or by calculating the sensitivity coefficient (Goyal, 2004; Lenhart et al., 2002; McKenney and Rosenberg, 1993). Because  $ET_o$  is a multi-factor parameter, its sensitivity varies with various CVs. Moreover, other factors, such as geological locations and the calculation models, also affect its sensitivity. Another way is to analyze the  $ET_o$  series re-estimated by the detrended CVs such as previous studies investigating the influences of detrended CVs on  $ET_o$  in the Changjiang River Catchment during 1970–2000 (Xu et al., 2006) and in northwestern China (Huo et al., 2013). However, their results only showed the influences of single CV on  $ET_o$ , not the simultaneous influences of multiple variables on  $ET_o$ . Because CVs change simultaneously, even the minimum, mean and maximum  $T$  ( $T_{\min}$ ,  $T_{\text{ave}}$  and  $T_{\max}$ ) increased with different patterns, not to mention the other CVs. It is not sufficient to only know the influences of single CV on  $ET_o$ . Previous research has not distinguished the ability of various  $T$ s ( $T_{\min}$ ,  $T_{\text{ave}}$  and  $T_{\max}$ ) on influencing  $ET_o$ . Moreover, not only removing linear trends from CVs is executable and referable, but also removing nonlinear trends in CVs is necessary because most CVs vary nonlinearly with time. There is limited research

that compares the influences of removing linear trends and nonlinear trends from various CVs on  $ET_o$ .

The effects of climate change on  $ET_o$  in northwestern China have been studied by Huo et al. (2013) using data obtained from 23 weather stations. The linear trends of  $T$ ,  $RH$ , solar radiation and  $U$  were removed, and the detrended CVs were used for re-estimating  $ET_o$ . Their results showed that the contribution of the decline in  $U$  to the decrease in  $ET_o$  is larger than that of other meteorological variables. This study investigates the effects of removing linear and nonlinear trends from the annual CVs (including  $T_{\min}$ ,  $T_{\text{ave}}$ ,  $T_{\max}$ ,  $RH$ ,  $U$  and  $n$ ) on the temporal variations of annual  $ET_o$  in Xinjiang, China. Our aims are to: i) assess the importance of each CV to  $ET_o$  for northern (NX), southern (SX) and the entirety of Xinjiang (EX); ii) to compare the differences in removing the trends from single, double and multi-CVs in various scenarios for each sub-region, and to distinguish the influences of various  $T$ -factors ( $T_{\min}$ ,  $T_{\text{ave}}$  and  $T_{\max}$ ) and non- $T$ -factors ( $U$ ,  $RH$  and  $n$ ) on changing  $ET_o$ ; and iii) to clearly show how the CVs affect  $ET_o$  in Xinjiang by comparing the original and recalculated  $ET_o$  series using the linear and nonlinear detrended CVs.

## 2. Data and methodology

### 2.1. The study sites and the data sets

The Xinjiang Uygur Autonomous Region is located in inland area of northwestern China. Xinjiang is surrounded on three sides by mountains and is distant from any sea (Li et al., 2016). The multiyear mean ratio of precipitation ( $P$ ) (147 mm) to  $ET_o$  (1512 mm) is 0.1 (Li and Sun, 2016); therefore, the region was classified as an arid zone according to Erinc (1965). A total of 53 weather stations in Xinjiang, China were selected as study sites (Fig. 1). The observed monthly weather data were collected from the Meteorological Data Sharing Service Network in China, with strict quality control. Elevations of the selected sites ranged between 30 and 3095 m. The data duration was from 1961 to 2013 with the completeness larger than 99.7%. Missing data were interpolated linearly with the neighboring months. The belt along Tuerqate – Aheqi – Bayinbuluke – Urumqi – Qitai – Qijiaoqing – Balikun – Zhuomaohu, which belongs to NX, was set as the border of NX and SX. There were 27 sites in NX and 26 sites in SX, respectively.

### 2.2. Estimation of reference crop evapotranspiration

The Penman-Monteith equation (Allen et al., 1998) is used here for estimating  $ET_o$  (mm):

$$ET_o = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{T_{\text{ave}} + 273} U_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (1)$$

where  $G$  is soil heat flux ( $\text{MJ m}^{-2} \text{day}^{-1}$ ),  $T_{\text{ave}}$  is for 2 m ( $^{\circ}\text{C}$ ),  $U_2$  is  $U$  at 2 m ( $\text{m/s}$ ),  $e_s$  is saturation vapor pressure (kPa),  $e_s - e_a$  is saturation vapor pressure deficit (kPa),  $\Delta$  is slope of vapor pressure curve ( $\text{kPa } ^{\circ}\text{C}^{-1}$ ),  $\gamma$  is psychrometric constant ( $\text{kPa } ^{\circ}\text{C}^{-1}$ ),  $R_n$  is net radiation ( $\text{MJ m}^{-2} \text{day}^{-1}$ ) and calculated with  $n$  following Allen et al. (1998). Monthly  $G$  is estimated by

$$G_M = 0.07(T_{M+1} - T_{M-1}) \quad (2)$$

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