



Effect of continental sources and sinks on the seasonal and latitudinal gradient of atmospheric carbon dioxide over East Asia



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HIGHLIGHTS

- Examination of regional scale atmospheric CO₂ column concentration over East Asia.
- The latest regional scale CO₂ simulation was conducted and evaluated.
- The recent ACOS-GOSAT, satellite CO₂ retrieval data was applied for the region.
- The characteristics of atmospheric CO₂ spatiotemporal variability over the East Asia.
- Identification of the main driving forces (sources/sinks) for the variabilities.

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ABSTRACT

Here we demonstrate the sharp seasonal and latitudinal gradient of atmospheric CO₂ over East Asia, where there are relatively few ground-based observations. The Greenhouse gases Observing SATellite (GOSAT) column-averaged dry air CO₂ mole fraction (xCO₂) retrieved by NASA's Atmospheric CO₂ Observations from Space (ACOS) (2009–2011) program and GEOS-Chem nested-grid CO₂ results are used. The strong anthropogenic emissions mainly from China and intensive vegetation uptake from northeastern Asia lead to a clear seasonal change of the xCO₂ between spring maximum and summer minimum (>10 ppm). In particular, the steep latitudinal gradient of summer time xCO₂ by 3–5 ppm in the vicinity of the Korean Peninsula (32°N–44°N) is likely attributed to the large difference in CO₂ fluxes among industry/cities, northeastern forests and the northwest Pacific region. This study represents the current progress to understand sub-continental scale atmospheric CO₂ variabilities with recent satellite retrievals and nested-grid modeling.

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1. Introduction

Quantitative understanding of the atmospheric carbon budget including biospheric carbon exchanges is crucial for climate mitigation policy in 21st century. In this aspect, there have been efforts to monitor global atmospheric carbon dioxide from space to overcome the spatial limits of ground-based observations (Buchwitz et al., 2007; Chahine et al., 2008; Crevoisier et al., 2009; Yokota et al., 2009; Kulawik et al., 2010). The Greenhouse gases

Observing Satellite (GOSAT) has observed carbon dioxide (CO₂) and methane (CH₄) column distributions since April 2009 (Maksyutov et al., 2008; Kuze et al., 2009; Yokota et al., 2009) and there have been continuous efforts to reduce the retrieval bias (Butz et al., 2011; Wunch et al., 2011; Yoshida et al., 2011; O'Dell et al., 2012). The GOSAT column CO₂ data has been recently applied to map the global CO₂ distributions with finer spatial and temporal resolution over land, but mapping CO₂ distribution solely by satellite measurements could not fully cover regional CO₂ variability mainly due to the limitation of the retrieval stems from cloud and aerosols (Hammerling et al., 2012). Here we focused on spatiotemporal variabilities of atmospheric CO₂ on a sub-continental scale. East Asia is the most human-dominated CO₂ emission region in the world, but there are relatively few monitoring sites, which restricts our understanding of spatial and seasonal CO₂ variability and relevant regional climate impacts. Under limited conditions,

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satellite observations such as Atmospheric CO₂ Observations from Space retrievals of the Greenhouse Gases Observing Satellite (ACOS-GOSAT) column averaged dry air mole fraction (xCO₂) and a 3-D chemical transport model simulation with nesting capability can be useful to determine the seasonal and spatial characteristics of the regional scale CO₂ variability over East Asia. In this study, we examined the seasonal and spatial characteristics of atmospheric CO₂ and tried to identify the main fluxes contributing the characteristics over East Asia (the region of 100°E–150°E, 20°N–50°N). The methods are explained in Section 2 and the spatiotemporal characteristics of regional CO₂ are explained in Section 3, followed by the model evaluation.

2. Methods

2.1. GOSAT CO₂ retrievals

GOSAT measures the concentrations of atmospheric carbon dioxide and methane with good sensitivity near the surface with the Thermal And Near-infrared Spectrometer for Carbon Observations (TANSO) Fourier Transform Spectrometer (FTS) observing in the Short Wave InfraRed (SWIR) region (0.76, 1.6, and 2.0 μm) and a wide Thermal Infrared (TIR) band (5.5–14.3 μm) with a circular 10.5 km FOV (Kuze et al., 2006, 2009). The satellite (also known as Ibuki) moves in the 666 km sun-synchronous orbit and completes one orbit in ~100 min providing global coverage in approximately 3 days (Kadygrov et al., 2009; Yokota et al., 2009). Recently, the bias of ACOS-GOSAT product v2.9 against the Total Carbon Column Observing Network (TCCON) observations has been greatly reduced mostly by correcting the bias stemming from the O₂ A-band (0.13 ± 1.97 ppm, Wunch et al. (2011); O'Dell et al. (2012)). We used ACOS-GOSAT column-averaged dry air mole CO₂ fraction products (xCO₂) (v2.9) from the NASA Goddard Earth Science Data and Information Services Center (<http://disc.sci.gsfc.nasa.gov/acdisc/data-holdings/acos-data-holdings>) from April 2009 to December 2011. We screened out the low-quality data with the master quality flag which defines retrieval success affected by cloud screening and retrieval algorithm diagnostics. The details of the retrieval method and data products are described elsewhere (Wunch et al., 2011; Crisp et al., 2012; O'Dell et al., 2012).

2.2. GEOS-Chem nested-grid CO₂ simulation for East Asia

We used GEOS-Chem, a global chemical transport model (v9-01-03, www.geos-chem.org) driven by GEOS-5 assimilated meteorology with nesting capability for East Asia. The GEOS-Chem CO₂ module was developed by Nassar et al. (2010), which updated the work by Suntharalingam et al. (2004). The GEOS-Chem model includes atmospheric CO₂ fluxes from biomass burning (year-specific Global Fire Emission Database version 3 (GFEDv3) (van der Werf et al., 2010)), biofuel burning (Yevich and Logan (2003)), fossil fuel combustion and cement production (from Carbon Dioxide Information and Analysis Centre (CDIAC) (Andres et al., 2011), ocean exchange (Takahashi et al. (2009)), international shipping (International Comprehensive Ocean-Atmosphere Data Set (ICOADS) (Corbett and Koehler, 2003, 2004), and aviation (Kim et al., 2007)). The fluxes from terrestrial biospheric exchange considered diurnal and seasonal variability by the Carnegie-Ames-Stanford-Approach (CASA) model (Potter et al., 1993) and residual annual exchange by inversion result from Baker et al. (2006). The CO₂ model also accounts for the chemical production of CO₂ from CO oxidation throughout the troposphere (Nassar et al., 2010). The main sources/sinks inventories such as fossil fuel emissions, biomass burning, and ocean exchanges are on a monthly basis. Further details of the model with sources/sinks are illustrated at Nassar et al. (2010).

The GEOS-Chem nested model uses native GEOS-5 meteorological fields with a horizontal resolution of 0.5° × 0.667° over Asia (70°E–150°E, 11°S–55°N, Chen et al., 2009). The lateral boundary conditions of CO₂ were archived from a GEOS-Chem global model run (2° × 2.5°) and used for every 3 h, consistent with the temporal resolution of the meteorological data. The details of the nested GEOS-Chem model are illustrated in Chen et al. (2009) and Wang et al. (2004). We initialized the model with a uniform global CO₂ concentration on January 1st in 2004 as 375 ppm, approximately equal to the global mean based on the measurements of marine surface mean CO₂ concentration (NOAA-ESRL-GMD sites) as previously set by Nassar et al. (2010), which was followed by a multi-year spin-up.

3. Results

3.1. The sharp seasonal and latitudinal CO₂ gradient over Korean Peninsula

The region for this study covers Eastern China, Korea, and Japan, the most populated region (110°E–150°E, 20°N–46°N) with the largest anthropogenic emissions in the world (Fig. 1). We find that the model simulation shows a steep latitudinal gradient of column-averaged CO₂ concentration around 40°N–45°N (will be shown in Fig. 3). In order to investigate that feature, we defined two zones, denoted “N” and “S”, which represent the northern zone (124°E–132°E, 40°N–44°N) and southern zone (124°E–132°E, 32°N–36°N) over the Korean Peninsula (Fig. 1). The ACOS-GOSAT xCO₂ time series of those zones from April 2009 to December 2011 is shown to demonstrate the seasonal and latitudinal gradient in Fig. 2. The red diamonds and black crosses represent xCO₂ of “N” and “S” zone respectively. In general, the xCO₂ reaches a maximum in April when vegetation is yet to start growing and reaches a minimum in July–August when photosynthesis is strongest over the region. In particular, the summer minimum of xCO₂ at “N” zone is as low as 385 ppm whereas that of southern zone is over 390 ppm in 2010 (Fig. 2). Even though highly cloudy condition during the Asian monsoon seriously restricts the availability of xCO₂, the clear difference of xCO₂ by 3–5 ppm between two zones is found in summer (June–August) (Fig. 2). The difference is larger than the level of standard deviation of the monthly mean (~2 ppm). The differences of GEOS-Chem column-averaged CO₂ results between two zones estimated from in July 2009 and 2010 are 3.6 and 3.5, respectively. That summertime latitudinal xCO₂ gradient over East Asia is substantially steep considering that the global model-simulated inter-hemispheric gradient of column CO₂ (45°N–45°S) by Olsen and Randerson (2004) is approximately 2.2 ppm. In addition, the latitudinal gradient of surface CO₂ over East Asia could be more significant since the north–south gradient of column CO₂ is about a half of the surface CO₂ based on the global analysis by Olsen and Randerson (2004).

The GEOS-Chem monthly mean column-averaged CO₂ in April (upper left in Fig. 3) and July (lower left) in 2009 can explain the significant seasonal difference. In April, there is higher column-averaged CO₂ (>391 ppm) moving through westerly transport from China (Sichuan province and Eastern China) through Korea and Japan, which is also evident from GEOS-Chem tagged CO₂ simulation for long-term contribution from fossil fuel and cement production (aggregated estimation since 2004, the panels at the second column). In order to estimate the seasonal contributions from fossil fuel emissions and biogenic exchange, we calculated the changes of contributions between January and April in 2009 (a part of non-growing months) and between April and July in 2009 (a part of growing months) from fossil fuel emissions (panels at the third column in Fig. 3) and biospheric exchange (panels at the fourth

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