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Numerical study of aircraft wake vortex evolution near ground in stable atmospheric boundary layer

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Abstract The evolutions of aircraft wake vortices near ground in stable atmospheric boundary layer are studied by Large Eddy Simulation (LES). The sensitivity of vortex evolution to the Monin-Obukhov (M-O) scale is studied for the first time. The results indicate that increasing stability leads to longer lifetimes of upwind vortices, while downwind vortices will decay faster due to a stronger crosswind shear under stable conditions. Based on these results, an empirical model of the vortex lifetime as a function of 10-m-high crosswind and the M-O scale is summarized. This model can provide an estimate of the upper boundary of the vortex lifetime according to the real-time crosswind and atmospheric stability. In addition, the lateral translation of vortices is also inspected. The results show that vortices can travel a furthest distance of 722 m in the currently-studied parameter range. This result is meaningful to safety analysis of airports that have parallel runways.

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19 1. Introduction

During an approach to an airport, the wake vortex of a leading aircraft poses danger to a following one. Therefore, it is necessary to study the wake vortex behavior in the atmospheric background. Fast-time wake prediction models have been developed by scholars.¹⁻⁴ In the early phase of the approach, the aircraft wake vortex is generated far from the ground,

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which is the so-called Out-of-Ground-Effect (OGE) condition. The wake vortex will turn into a rapid decay in tens of seconds as a result of the development of linear instability^{5,6} due to the multi-reaction of the vortex pair. Furthermore, the vortex pair will leave the flight path by lateral displacement due to the crosswind, or by descending under the induction of each other, both of which contribute to the safety of the following aircraft. 26

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In the final approach, the wake vortex is generated near the ground. Compared to the OGE condition, the In Ground Effect (IGE) condition requires more concerns. In the ground effect, image vortices will increase the separation of the vortex pair,⁷ which weakens the interaction between the vortex pair and decreases the vortex decay rate. Besides, the vortices will rebound in the ground effect⁷ and come back to the glide slope. These factors increase the wake-encounter risk for the following aircraft.

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42 One of the early IGE wake prediction models was developed by Proctor et al.⁸ They revealed that the vortices turn 43 into a rapid decay soon after the rebound from the ground. 44 In addition, the evolution of vortex circulation after the 45 rebound is described with an exponential model based on 46 Large Eddy Simulation (LES) results. Except for the circula-47 48 tion, the displacement of the vortices was modeled by Robins et al.⁷ In their work, the separation of the vortex pair is mod-49 eled by bringing in image vortices below the ground, and the 50 rebound is modeled by introducing secondary vortices. The 51 influence of the atmospheric background was taken into con-52 sideration by Holzäpfel and Steen.⁹ In their model, the decay 53 54 rate of vortex circulation is related to the 10-m-height crosswind velocity. The newest model is the Deterministic wake 55 Vortex Model (DVM) developed by De Visscher et al. in 56 57 2013¹⁰, in which Holzäpfel's model⁹ is used to predict the vor-58 tex circulation, and the vortex displacement is predicted by the 59 method of discrete vortices, in which vortex particles are 60 released from the ground to simulate ground induced secondary vortices.^{11,1} 61

However, the influence of atmospheric stability has not 62 been considered in the aforementioned near-ground models. 63 In an unstable boundary layer (or convective boundary layer), 64 vortices will deform and decay rapidly under updraft and 65 downdraft velocities, which is to the benefit of the following 66 aircraft's safety.13 The influence of a stable boundary layer will 67 68 be much more complex. On one hand, the stable stratification 69 can suppress the atmospheric turbulence, which will lead to a longer wake vortex lifetime. On the other hand, the stratifica-70 tion will promote the vortex decay by buoyancy effects.^{14–16} 71 Therefore, it is necessary to study the wake vortex decay in 72 73 the stable atmospheric boundary layer to improve wake pre-74 diction models.

75 In the current OGE wake vortex prediction models, the B runt-Väisälä (BV) frequency and the Eddy Dissipation Rate 76 77 (EDR) are used as the parameters describing stability and atmospheric turbulence, respectively.¹⁻³ In the IGE condition, 78 these two parameters are not enough because wind shear and 79 80 ambient turbulence kinetic energy can also influence the wake vortex evolution. However, it is impractical to bring all these 81 82 atmospheric parameters into a fast prediction model because it is difficult to establish an empirical model with so many 83 independent variables. In this paper, the Monin-Obukhov 84 (M-O) scale L^{17} is chosen as the parameter to measure the 85 atmospheric stability. On one hand, in the IGE condition, vor-86 tices are generated in the atmospheric surface layer (below 87 60 m), where the BV frequency and the EDR are not mutually 88 independent and can be related to the M-O scale.¹⁷ According 89 to the M-O simulation theory, atmospheric information 90 including wind and potential temperature profile and the ambi-91 ent turbulence can be estimated from the 10-meter-height 92 crosswind V_{10} and the M-O scale L. In this way, the number 93 94 of independent atmospheric parameters is prominently 95 reduced. On the other hand, a model using the M-O scale 96 may be more practical because the M-O scale can be easily estimated from the wind speed and the net radiation index^{18,19} 97 which can be provided by the airport control tower without 98 additional observing facilities. In this paper, LES is applied 99 to the wake vortex decay process in the stable atmospheric 100 boundary layer. The vortex evolution is studied under different 101 V_{10} and L, and simulation results are discussed and modeled. 102 The paper is organized as follows. The numerical method 103

and the physical model are introduced in Section 2. The results are displayed and discussed in Section 3. Finally, main conclusions are summarized in Section 4.

2. Methodology

2.1. Numerical methods

The LES in this paper solves the Boussinesq-approximated Navier-Stokes equations.

$$\frac{\partial u_i}{\partial x_i} = 0 \tag{1}$$

$$\frac{\partial u_i}{\partial t} + u_j \frac{\partial u_i}{\partial x_j} = -\frac{1}{\rho} \cdot \frac{\partial p}{\partial x_i} + \frac{\partial}{\partial x_j} \left[(v + v_t) \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right) \right]$$
(2)

$$\frac{\partial\theta}{\partial t} + u_j \frac{\partial\theta}{\partial x_j} = \frac{\partial}{\partial x_j} \left[Pr_t(v + v_t) \frac{\partial\theta}{\partial x_j} \right]$$
(3)

where u_i and u_i are the velocity component and x_i and x_i are the coordinate, with p the pressure, ρ the density, t the time. v and v_t are the molecular kinematic viscosity and sub-grid scale viscosity, respectively. In the equation of the potential temperature θ , the turbulent Prandtl number Pr_t is set to 0.72. The Lagrangian dynamic sub-grid model²⁰ is used.

These equations are solved numerically on a movable, non-126 staggered Cartesian grid which is automatically adapted to the 127 real-time flow field to obtain a higher resolution in a large gra-128 dient region such as the vortex cores with a limited total grid 129 number. The self-adaptive algorithm is based on Gnoffo and 130 Nakahashi's spring analogy method.^{21,22} The Finite Volume 131 Method (FVM) is applied for the spatial discretization and a 132 fourth-order Runge-Kutta integration is used in the time 133 advancement. The details of the numerical method are 134 described in Refs.^{23,24} This method has been proven reliable 135 in both wake vortex simulation¹⁸ and micro-scale atmospheric 136 boundary layer LES^{25,26} 137

2.2. Setup of the initial field 138

The initial field is combined with the atmospheric background 139 and a pair of wake vortices. In the background atmosphere, 140 the profile of wind V and the potential temperature θ in the 141 surface layer are initialized with the M-O simulation theory² 142 143 as

$$V = \frac{v_{r}}{\kappa} \left(\ln \frac{z}{z_{0}} + 5\frac{z}{L} \right) 0 \leqslant z < L$$

$$V = \frac{v_{r}}{\kappa} \left(\ln \frac{z}{z_{0}} + 4 + 5 \ln \frac{z}{L} + \frac{z}{L} \right) L \leqslant z < 2L$$

$$L = \frac{\theta_{0}v_{r}^{2}}{\kappa g \theta_{r}}$$
(4)

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$$\theta = \frac{\theta_r}{\kappa} \left(\ln \frac{z}{z_0} + 5\frac{z}{L} \right) 0 \leqslant z < L$$

$$\theta = \frac{\theta_r}{\kappa} \left(\ln \frac{z}{z_0} + 4 + 5\ln \frac{z}{L} + \frac{z}{L} \right) L \leqslant z < 2L$$
(5)

where v_{τ} and θ_{τ} are the friction velocity and friction tempera-150 ture, respectively. L is the M-O scale and $\theta_0 = 290$ K is the ref-151 erence temperature. $\kappa = 0.4$ is the Karman constant, and 152 $g = 9.8 \text{ m/s}^2$ is the gravity acceleration. z_0 is the roughness 153 Download English Version:

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