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Optimal convergence rates of the supercritical surface quasi-geostrophic equation

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ABSTRACT

This study is concerned with the optimal convergence rates of the supercritical surface quasi-geostrophic equation. When the global weak solution  $\theta$  of the supercritical surface quasi-geostrophic equation satisfies

$$\nabla\theta \in L^r(0, \infty; \dot{B}_{p,\infty}^0(\mathbb{R}^2)) \quad \text{for } \frac{2}{p} + \frac{\alpha}{r} = \alpha, \quad \frac{2}{\alpha} < p < \infty,$$

then even for large initial data perturbation, every weak solution  $\tilde{\theta}(x, t)$  of the perturbed quasi-geostrophic equation converges to  $\theta(x, t)$  with the optimal algebraic convergence rate

$$\|\tilde{\theta}(t) - \theta(t)\|_{L^2} \leq C(1+t)^{-\frac{1}{\alpha}} \quad t > 0.$$

The findings are mainly based on some new estimates for the trilinear form in Besov spaces and the generalized Fourier splitting method.

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1. Introduction

In this study, we consider the Cauchy problem of the supercritical dissipative surface quasi-geostrophic fluid motion model given by Constantin, Majda and Tabak [1]

$$\begin{cases} \left(\frac{\partial}{\partial t} + u \cdot \nabla\right)\theta + \kappa\Lambda^\alpha\theta = f, & (x, t) \in \mathbb{R}^2 \times (0, \infty), \\ \theta(x, 0) = \theta_0, & x \in \mathbb{R}^2. \end{cases} \quad (1.1)$$

Here  $0 < \alpha < 1$  and  $\kappa > 0$  is a dissipative coefficient.  $\Lambda$  is the Riesz potential operator defined by  $\Lambda = (-\Delta)^{1/2}$ ,  $\theta(x, t)$  is an unknown scalar function representing potential temperature.  $u(x, t)$  is the velocity field determined by

$$u = -\nabla^\perp(-\Delta)^{-\frac{1}{2}}\theta = -\mathcal{R}^\perp\theta = (\mathcal{R}_2\theta, -\mathcal{R}_1\theta) \quad (1.2)$$

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where  $\mathcal{R}_j, j = 1, 2$ , is the 2D Riesz transform.

Due to its mathematical importance and its potential for applications in atmospheric and oceanographic sciences [2], on one hand, there is a large literature on the well-posedness for solutions of the surface quasi-geostrophic equation. When  $\alpha = 1$ , the surface quasi-geostrophic equation (1.1) shares similar features with the three-dimensional incompressible Navier–Stokes equations. Thus  $\alpha = 1$  is therefore referred as the critical case, while the cases  $0 < \alpha < 1$  and  $1 < \alpha \leq 2$  are supercritical and subcritical, respectively. In the critical and subcritical cases, the global smooth solution for the sufficiently smooth initial data have been proved by Constantin and Wu [3], Caffarelli and Vasseur [4], Kiselev, Nazarov and Volberg [5], Constantin and Vicol [6], respectively. However, the question of global existence for smooth solution of the supercritical quasi-geostrophic equation is still unanswered. One may refer to many important studies on global solutions in some critical spaces or regularity criteria of weak solutions have been examined (see [7–12] and references therein).

On the other hand, it is desirable to understand the asymptotic behavior of the surface quasi-geostrophic equation (1.1). By using the classic Fourier splitting methods, Constantin and Wu [3] firstly proved the  $L^2$  decay of weak solution of zero-forced quasi-geostrophic equation (1.1)

$$\|\theta(t)\|_{L^2} \leq C(1+t)^{-\frac{1}{\alpha}}. \tag{1.3}$$

The above result essentially shows there exists a global weak solution of (1.1) that converges to the trivial solution  $\tilde{\theta} \equiv 0$  with the rate  $(1+t)^{-\frac{1}{\alpha}}$ . The result is further studied by many authors such as Dong and Liu [13], Carrillo [14], Ferreira, Niche and Planas [15], Niche and Schonbek [16], Zhou [17] and references therein.

When  $f \neq 0$  in (1.1), it is clear that (1.1) has no any constant solutions. It is natural and important to consider the asymptotic convergence issue of nontrivial solutions of the nonzero forced surface quasi-geostrophic equation (1.1). More precisely, we consider the following perturbed surface quasi-geostrophic equation

$$\begin{cases} \left(\frac{\partial}{\partial t} + v \cdot \nabla\right) \tilde{\theta} + \kappa \Lambda^\alpha \tilde{\theta} = f, \\ v = -\nabla^\perp (-\Delta)^{-\frac{1}{2}} \tilde{\theta} = -\mathcal{R}^\perp \tilde{\theta} = (\mathcal{R}_2 \tilde{\theta}, -\mathcal{R}_1 \tilde{\theta}), \\ \tilde{\theta}(x, 0) = \theta_0 + w_0 \end{cases} \tag{1.4}$$

where  $w_0(x)$  is the initial data perturbation. On one hand, as for the uniform convergence, Chae and Lee [18] proved the global solution  $\theta$  of the critical surface quasi-geostrophic equation converges to  $\tilde{\theta}$  in Besov space  $C([0, \infty); B_{2,1}^1) \cap L^1((0, \infty); \dot{B}_{2,1}^2)$  under the small initial perturbation. Dong and Chen [19] proved the global solution  $\theta$  of the critical and supercritical surface quasi-geostrophic equation asymptotically converges to  $\tilde{\theta}$  in critical Lebesgue space

$$\nabla \theta \in L^r(0, \infty; L^p(\mathbb{R}^2)), \quad \frac{2}{p} + \frac{\alpha}{r} = \alpha, \quad \text{and} \quad \frac{2}{\alpha} < p < \infty$$

for arbitrary initial data and external force perturbations. Ren and Ma [20] studied the global solution  $\theta$  of the subcritical surface quasi-geostrophic equation asymptotically converges to  $\tilde{\theta}$  in Lorentz space. Jia, Gui and Dong [21] further examined the asymptotic stability of the supercritical surface quasi-geostrophic equation in critical BMO spaces. The first and the third authors [22] recently examined the asymptotic convergence rates of the supercritical surface quasi-geostrophic equation in the critical Morrey space which is larger than Lorentz space and Lebesgue space. We would like to mention that Dai [23] proved the global solution  $\theta$  of the subcritical surface quasi-geostrophic equation asymptotically converges to the solution of the following steady quasi-geostrophic equation

$$\begin{cases} (v \cdot \nabla) \tilde{\theta} + \kappa \Lambda^\alpha \tilde{\theta} = f, \\ v = -\nabla^\perp (-\Delta)^{-\frac{1}{2}} \tilde{\theta} = -\mathcal{R}^\perp \tilde{\theta} = (\mathcal{R}_2 \tilde{\theta}, -\mathcal{R}_1 \tilde{\theta}) \end{cases} \tag{1.5}$$

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