



Fluvial history of the Sub-Carpathian Basins (Poland) during the last cold stage (60–8 cal ka BP)



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ABSTRACT

The fluvial landforms of the last cold stage are represented in the Sub-Carpathian Basins by a system of 2–3 terraces, which grade into alluvial fans of a height of 15–20, 8–12 and 6–8 m above the present riverbeds in the mountain foreland. In the Vistula River valley the 15–20 m high terrace, covered with loess, extends along the western margin of the Sandomierz Basin and northern margin of the Carpathians. Away from the mountain margin, it passes into the sandy terrace plain with dunes. It is formed of at least two Interpleniglacial alluvial fills: the older one dated at 48–36 ka BP (over 50–40.6 ka cal BP) and the younger one dated 30–25 ka BP (34.1–29.0 ka cal BP) thus pertaining to the Upper Pleniglacial. This latter forms a terrace 15 m high. In general, accumulation of fluvial sediments interrupted with erosion phases dominated in the mountain foreland during the period 60–25 ka BP (over 60–29 ka cal BP). The most significant erosion phase took place before the maximum extension of the Vistulian (Weichselian, Wisconsinian) ice sheet (i.e. before 25–20 ka BP = 29–24 ka cal BP) and was connected with a change from an oceanic to a more continental climate. The next alluvial fill of terrace 8–12 m high (with remnants of braided river channels) was formed at the end of the Upper Plenivistulian. The incision of the river channels below the present channel level was followed by aggradation in the Late Vistulian (15–13 ka BP = 18.2–15.6 ka cal BP) accompanied by a change in river channel system from braided to meandering. In the Wistoka and San River valleys alluvial plains extending by lateral erosion of large meanders were abandoned by rivers mainly in the Younger Dryas and the Preboreal. The Late Vistulian large palaeomeanders are missing along lower Soła and lower Dunajec. These rivers, fed from higher mountains, had still tendency to braiding. The climate cooling of the Younger Dryas brought an increase in the delivery of bedload to the river channels and a tendency to braiding in Vistula valley (Drwinka depression). The expansion of forest in the Early Holocene reduced the water discharge and the size of meanders. As a result, their width fallen by 3–5 times. Alluvial data showed also a distinct humid climatic phase ca. 9.5–8.5 ka cal BP.

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1. Introduction

The Sub-Carpathian Basins represent a depositional basin gathering the sediments of rivers flowing from the Carpathians. The Vistulian (Weichselian, Wisconsinian) was a stage which saw the intensive accumulation of alluvia in this area. During Last Glacial Period extensive alluvial fans in the outlet of Carpathian valleys to the mountain foreland were formed. Alluvial fans were

dissected by river incision and in the widened trough covers of younger alluvial fans were inserted. This resulted in the formation of alluvial terraces. The best recorded depositional sequence represents the Middle and the Upper Plenivistulian, correlated with a large part of MIS 3 and MIS 2 (Dansgaard et al., 1993). Owing to the widespread use of the radiocarbon method and luminescence methods (TL and OSL), the stratigraphic–palaeogeographic evolution of the Sub-Carpathian Basins during the period 60–8 cal ka BP, studied as part of the scientific project INTIMATE (COST Action ES0907, Blockley et al., 2012), is among the best recorded in Europe. Using these dating methods allows a reconstruction of changes in fluvial deposition along the valleys situated in the

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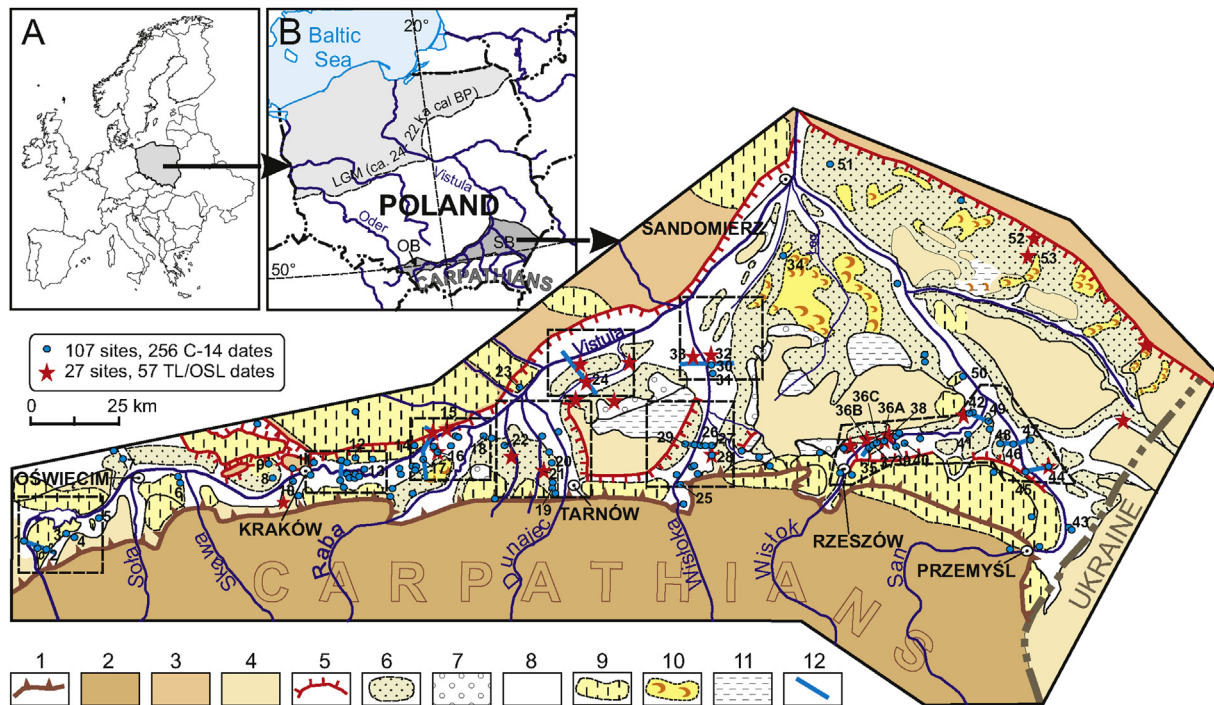


Fig. 1. A geomorphological map of the Fore-Carpathian Basins and the distribution of the most important sites of fluvial deposits dated by the radiocarbon and thermoluminescence (TL, OSL) methods (after Starkel, 1980; Gębica, 2004, partly modified). 1 – margin of the Carpathians, 2 – Carpathian foothills, 3 – Uplands, 4 – Fore-Carpathian plateaus, 5 – main escarpments, 6 – Vistulian terraces and alluvial fans, 7 – older (than Vistulian) river terraces, 8 – Holocene valley floors (floodplains), 9 – loess cover on the terraces and plateaus, 10 – fields of drift-sand and dunes (inactive), 11 – slope-wash plains and cones, 12 – geological cross-sections. The most important sites are numbered (1–53). Sites dated by radiocarbon method are marked with circles, and by luminescence method with stars. Locations of valley sections described in the text and presented in Figs. 2–10 are marked by dashed rectangles. List of the most important sites (mentioned in the text): 1 – Drogomyśl, 2 – Pierściec, 3 – Chybie, 4 – Bronów, 5 – Kaniów, 6 – Zator, 7 – Bieruń Nowy, 8 – Szejowice, 9 – Kryspinów, 10 – Ludwinów, 11 – Prądnik alluvial fan in Kraków, 12 – Nowa Huta, 13 – Pleszów, 14 – Brzesko Nowe, 15 – Śmitowice, 16 – Trawniki, 17 – Drwinka–Grobla Forest, 18 – Gróbka, 19 – Dębina Łętowska, 20 – Szujec, 21 – Niwka near Radłów, 22 – Włoszyn, 23 – Opatowiec, 24 – Zabrze–Czapłówka, 25 – Podgrodzie, 26 – Brzeźnica B, 27 – Brzeźnica A–C, 28 – Pustynia, 29 – Wola Żyrakowska, 30 – Mielec-Trzeźń Mała, 31 – Mielec-Wojślawskie depression, 32 – Chorzelów, 33 – Trzciana, 34 – Kobyłarnia, 35 – Łąka, 36 – Łukawiec I–IV, 37 – Łukawiec V, 38 – Czarna-Podbór, 39 – Wola Mała 1, 40 – Wola Mała 2, 41 – Gniewczyzna Łańcucka, 42 – Grodzisko Nowe, 43 – Stubno, 44 – Bobrówka-Wietlin, 45 – Jarosław, 46 – Manasterz-Nieplekowiec, 47 – Kostków, 48 – Wólka Petkińska, 49 – Wólka Ogrzykowska, 50 – Leżajsk, 51 – Łązek, 52 – Kąty, 53 – Korytków.

periglacial zone (Fig. 1) (Rotnicki, 1987; Harasimiuk, 1991; Starkel, 1991, 1994, 1995b; Starkel and Gębica, 1995; Turkowska, 1995; Superson, 1996; Gębica, 2004; Starkel et al., 2007; Gębica, 2013; Starkel et al., 2015).

The aim of this paper is a reconstruction of the chronostratigraphy of alluvia and a determination of the age of the Vistulian alluvial terraces on the base of large data-set. A leading point is the

timing of the phases of alluvial deposition and valley incision in response to changing environmental conditions in the river catchments. The authors also attempted to correlate the terrace surfaces, differentiated in structure and morphology, along the longitudinal transects of valleys. This review paper presents all collected dating results published originally mainly in local literature, or in press (Tables 1 and 2).

Table 1
Summary of dating results for the most important 71 sites (marked by numbers (53) and numbers with letters for sub-localities) from the Sub-Carpathian Basins. All radiocarbon dates were calibrated using OxCal 4.2 programme (Bronk Ramsey, 2009) and IntCal13 calibration curve (Reimer et al., 2013). Results of calibration were rounded by 100 yrs. Sites are organized along particular river valleys. Sites from Vistula valley are interrupted by sites from its tributaries (from W to E). For location of sites see Fig. 1. Dates for samples from alluvial fans, used for analysis presented in Fig. 11 are marked by “***”. Dates for samples from palaeochannels, used for analysis presented in Fig. 13 are marked by “****”.

Name Altitude and coordinates	No. at the map	River Valley	Terrace level	Dated material	Lab. Co.	Dating method	Date	Error	Calibrated age cal BP (68.2% conf. Int.)	References
Drogomyśl 262 m a.s.l. 18.75 E, 49.867 N	1	Vistula	IV	peat peat peat peat organic clay organic silt	Lv-1074 Gd-1393 Lv-966 Lv-965** Gd-1395** Gd-913	C-14 C-14 C-14 C-14 C-14 C-14	9020 9490 11220 11250 31100 30000	120 100 100 110 1100 +5000 –3000	10300–9900 11100–10500 13300–12900 13300–13000 36300–34000 42300–31100	Niedziałkowska et al. (1985)
Pierściec I 270–280 m a.s.l. 18.816 E, 49.835 N	2	Vistula	I	peat organic silt peat organic silt	Gd-6362 Gd-5784 Gd-6364** Gd-6246**	C-14 C-14 C-14 C-14	29200 29500 30700 35800	1000 800 1300 2500	34300–32000 34500–32700 36400–33500 43100–37600	Niedziałkowska and Szczepanek (1993–1994)
Chybie I 260 m a.s.l. 18.828 E, 49.902 N	3A	Vistula	I	peat peat	Gd-948 Gd-1460	C-14 C-14	>41000 >45000	>44500 >48400		Niedziałkowska et al. (1985)

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