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Spatial and historical distribution of organic phosphorus driven by environment conditions in lake sediments

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ABSTRACT

The chemistry of sedimentary organic phosphorus (OP) and its fraction distribution in 16 sediments are greatly influenced by environmental conditions such as terrestrial inputs 17 and runoffs. The linkage of OP with environmental conditions was analyzed on the basis of 18 OP spatial and historical distributions in lake sediments. The redundancy analysis and OP 19 spatial distribution results suggested that both NaOH-OP (extracted by NaOH-OP) and Re-OP 20 (residual OP) in surface sediments from the selected 13 lakes reflected the gradient effects of 21 environmental conditions and the autochthonous and/or allochthonous inputs driven by 22 latitude zonality in China. The lake level and salinity of Lake Hulun and the runoff and 23 precipitation of its drainage basin were reconstructed on the basis of the geochemistry 24 index. This work showed that a gradient in weather conditions presented by the latitude 25 zonality in China impacts the OP accumulation through multiple drivers and in many ways. 26 The drivers are mainly precipitation and temperature, governing organic matter (OM) 27 production, degradation rate and transportation in the watershed. Over a long temporal 28 dimension (4000 years), the vertical distributions of Re-OP and NaOH-OP based on a dated 29 sediment profile from HLH were largely regulated by the autochthonous and/or allochtho- 30 nous inputs, which depended on the environmental and climate conditions and 31 anthropogenic activities in the drainage basin. This work provides useful environmental 32 geochemistry information to understand the inherent linkage of OP fractionation with 33 environmental conditions and lake evolution. 34

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49 Introduction

In the last decade, the importance of organic phosphorus (OP) has been gradually recognized as a potentially large pool of bioavailable P due to the fact that the OP fraction is at least as abundant as inorganic P in many waters and sediments (Turner et al., 2005). Therefore, many publications are available concerning OP decomposition (Dyhrman et al., 2006; Stets and Cotner, 2008), fractionation (Ivanoff et al., 1998a; Turner et al., 56 2005; Zhu et al., 2013) and OP compound determination using 57 ³¹P nuclear magnetic resonance (NMR) spectroscopy (Ahlgren 58 et al., 2005; Turner and Leytem, 2004), which has greatly 59 elevated the acknowledged abundance of the OP fraction in 60 the environment. However, knowledge gaps still exist regarding 61 the factors driving the underlying processes governing the 62 spatial and historical distribution of OP pools in lake sediments. 63

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Lakes generally have high rates of sedimentation and primary production holding information on organic carbon (OC), nitrogen and phosphorus, so short-term processes that affect organic matter (OM) delivery and burial are amplified in lake sediment records (Meyers, 2003), preserving valuable in situ historical information about the past condition of lakes

and their environmental states (Zhang and Shan, 2008). 70 Therefore, vertical concentration profiles of total OC (TOC) 71 72archived in lacustrine sediments are widely used as a tool to 73 reconstruct the paleoenvironment and paleoclimate of the lakes' basins (Lü et al., 2008; Lüniger and Schwark, 2002; 74 Meyers, 2003; Meyers and Ishiwatari, 1993; Wu et al., 2007) and 75monitor the effects of environment, climate change and/or 76 anthropogenic activities on the environment (Bränvall et al., 77 2001; Herczeg et al., 2001; Meyers, 2003; Renberg et al., 2002; 78 79 Routh et al., 2004). Usually, significant correlations are found between natural OM and OP, because OP is one of the main 80 components of OM, and OM acts as the main carrier for OP in 81 the environment. Accordingly, the chemistry of sedimentary 82 OP and the distribution of its forms in sediments is also 83 greatly influenced by environmental and climate conditions 84 such as terrestrial inputs and runoffs. The geochemical 85 phases of OP are important parameters in understanding the 86 87 burial, diagenesis, bioavailability and environmental geo-88 chemical significance of OP in sediments.

89 As mentioned above, previous works mainly concern the 90 OP abundance and the determination of its compounds in the 91 surface sediments/soils, and some works are beginning to focus the distribution, degradation, diagenesis, transforma-92 tion and their factors for OP in sediment profiles using 93 fractionation techniques and/or ³¹P NMR (Ahlgren et al., 94 2006; Coelho et al., 2004; Ding et al., 2013; Edlund and 95 Carman, 2001; Penn et al., 1995; Reitzel et al., 2007; Xu et al., 96 2013; Zhang et al., 2013) due to the environmental significance 97 of OP. In an attempt to investigate the relationship between 98 OP fractions in sediments and environmental conditions, 99 this work hypothesized that the trends of OP fractions in 100 surface sediments, which were collected from different lakes 101 spanning a wide latitude range in China (spatial distribution), 102 were driven by the environmental conditions of the lake 103basins, which were next in importance to latitude effects. 104 105Then the corresponding relationship of OP fractions to environmental conditions and lake evolution over a long time 106 span (4000 years) was investigated based on a dated sediment 107profile from Lake Hulun (HLH). Accordingly, the objectives 108 of this study were to (1) present the trends and factors of OP 109 fractionation in surface sediments from 13 lakes on the basis 110 of redundancy analysis and spatial distribution, (2) recon-111 struct the lake level and salinity and the runoff and 112 precipitation of the lake basin, and discuss the relationship 113 of OP fractionation with environmental conditions and 114 climate changes according to the historical distribution of OP 115 fractions in a dated sediment profile, (3) reveal the main 116 processes or mechanism driving OP fractionation in the 117 118 aquatic environment. This work provides some useful information on OP fractions archived in sediments and shares 119 the novel concept and results with the environmental science 120community, especially the response of OP fractionation to 121 environmental conditions over such large spatial and tempo-122 ral dimensions. 123

1. Materials and methods

1.1. Study area

Lake Hulun (HLH) is located in northeastern Inner Mongolia, 127 China. This is at the northeastern corner of the monsoon 128 margin, and thus is sensitive to small variations in the East 129 Asian monsoon (An, 2000; Wen et al., 2010). The highly 130 variable climate is generally arid with an annual mean 131 precipitation of 290 mm, of which three-fifths falls between 132 July and August, an annual mean theoretical evaporation of 133 1600 mm, and an annual mean air temperature of -0.2°C. 134 It once was the fifth largest lake in China, with an area 135 of 2339 km² and an average depth of 5.7 m. During the last 136 35-50 years the lake has experienced a dramatic reduction 137 in its size and water depth due to variations in climate. In 138 2011, its area had shrunk to 1770 km² with the mean depth of 139 3.0-3.5 m. Two major permanent rivers, the Herlun (Kerulen 140 or Kelulun) and Urshen (Orxon or Orshun) River, feed the 141 lake from the southwest and east, respectively (Fig. 1). The 142 Dalanolom River, connected to the lake in the northeast, used 143 to be an intermittent river. Since 1971 this river has been a 144 channel used for water exchange between Lake Hulun and 145 the Hailaer River. For much of the year, however, the water is 146 stagnant, and occasionally, when its water table above sea 147 level is lower than 545.1 m, there is a reverse flow of waters 148 into the lake (Sun, 2010). 149

1.2. Sediment sampling

Sediments from HLH were sampled in August 2011. The 151 surface sediments (0-10 cm) were collected using a KC col- 152 lector mod B (Swedaq), while sediment cores were sampled 153 applying a SA Beeker collector (Eijkelkamp). The sediment 154 cores (HLH15, Fig. 1) were sectioned into 2-cm slices immedi- 155 ately after collection and stored in sealed polyethylene bags. 156 Samples were transported to the laboratory in iceboxes (<4°C). 157 At the laboratory the sediments were freeze-dried, ground 158 gently in a mortar and sieved through a standard 100-mesh 159 sieve prior to analysis. In addition, the physiochemical data of 160 sediments that were sampled in a similar manner from Lake 161 Wuliangsuhai (WLSH) are included in this study, which is 162 detailed in previous work (Lü et al., 2016a). 163

1.3. Experiments and methods

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A sequential extraction procedure modified by Ivanoff et al. 165 (1998a) and Zhang et al. (2008) was used to fractionate the OP 166 pools in the surface sediments and sediment cores. Following 167 the procedure, OP in sediments is fractionated into a labile 168 pool, a moderately labile pool, and a nonlabile pool. The labile 169 pool is extracted with 0.5 mol/L NaHCO₃ at pH 8.5. The 170 extracted P includes both OP (LOP) and inorganic P (IP) in 171 pore water and sorbed on sediments. The moderately labile 172 pool is extracted with 1.0 mol/L HCl (including HCl-OP, and IP), 173 followed by 0.5 mol/L NaOH. The NaOH extract is acidified 174 with concentrated HCl to separate the nonlabile fraction 175 (humic acid fraction, HAOP) from the moderately labile fraction 176 (fulvic acid fraction, FAOP). Finally, the highly resistant, nonlabile 177

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