Accepted Manuscript

Scavenging ratio of black carbon in the Arctic and the Antarctic

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PII: S1873-9652(17)30118-4

DOI: 10.1016/j.polar.2018.03.002

Reference: POLAR 377

To appear in: Polar Science

Received Date: 23 October 2017

Revised Date: 15 March 2018

Accepted Date: 16 March 2018

Please cite this article as: Gogoi, M.M., Babu, S.S., Pandey, S.K., Nair, V.S., Vaishya, A., Girach, A.I., Koushik, N., Scavenging ratio of black carbon in the Arctic and the Antarctic, *Polar Science* (2018), doi: 10.1016/j.polar.2018.03.002.

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1 Scavenging Ratio of Black Carbon in the Arctic and the Antarctic

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7 Abstract

8 Long-term monitoring of atmospheric aerosols and their interaction with radiation, cloud, and cryosphere over the Arctic and the Antarctic are very important for the global climate change 9 10 related issues. In this regard, for conducting aerosol measurements, India has extended the concerted efforts to the Indian stations at the Svalbard region of the Norwegian Arctic (Himadri, 11 78°55'N 11°56'E, 8 m a.s.l.) in the northern hemisphere and the Larsemann Hills of coastal 12 Antarctic (Bharati, 69°24.4'S 76°11.7'E, 40 m a.s.l.) in the southern hemisphere. In the present 13 study, we have examined darkening of snow due to black carbon (BC) deposition in different 14 sunlit seasons and estimated the scavenging ratio of BC over both the poles from simultaneous 15 measurements of atmospheric BC and BC deposited in snow. The study reveals distinct spatio-16 temporal variability of BC in polar snow, even though the magnitude of BC concentrations is, in 17 general, low (< 12 ppbw, parts per billion by weight). During local summer seasons, the 18 concentration of BC in Arctic snow (median ~ 7.98 ppbw) was higher compared to the values at 19 Antarctica (median ~ 1.70 ppbw). Concurrent with this, the scavenging ratio (SR) also showed 20 large variability over both the poles. Relatively higher values of SR over the Antarctica (mean ~ 21 119.54 ± 23.04 ; during southern hemispheric summer) in comparison to that over the Arctic 22 (mean ~ 69.48 \pm 4.79; during northern hemispheric spring) clearly indicates the difference in 23 removal mechanisms (aerosol mixing, aging and size distribution) of BC from the atmosphere 24 over distinct polar environments. Measurement of spectral incoming and reflected radiances over 25 the Arctic snow during the early spring season of 2017 indicated the values of surface broadband 26 albedo varying between 0.64 and 0.79. The Snow, Ice and Aerosol Radiative (SNICAR) model 27 simulated values of spectral albedo correlated well with the measured ones and indicated the role 28 29 of dust absorption, in addition to that of BC, in changing the snow albedo. This information needs to be accurately incorporated in the radiative transfer models for the accurate estimation of 30 snow albedo forcing over the Polar Regions. 31

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