

Available online at www.sciencedirect.com

J. Math. Anal. Appl. 304 (2005) 783–796

Journal of MATHEMATICAL ANALYSIS AND **APPLICATIONS**

www.elsevier.com/locate/jmaa

Bifurcating periodic solutions of wind-driven circulation equations

Zhi-Min Chen ^a*,*b*,*[∗] , W.G. Price ^a

^a *School of Engineering Sciences, Ship Science, University of Southampton, Southampton SO17 1BJ, UK* ^b *School of Mathematics, Nankai University, Tianjin 300071, People's Republic of China*

> Received 5 June 2004 Available online 28 January 2005

> > Submitted by C.E. Wayne

Abstract

The existence of bifurcating periodic flows in a quasi-geostrophic mathematical model of winddriven circulation is investigated. In the model, the Ekman number *r* and Reynolds number *R* control the stability of the motion of the fluid. Through rigorous analysis it is proved that when the basic steady-state solution is independent of the Ekman number, then a spectral simplicity condition is sufficient to ensure the existence of periodic solutions branching off the basic steady-state solution as the Ekman number varies across its critical value for constant Reynolds number. When the basic solution is a function of Ekman number, an additional condition is required to ensure periodic solutions.

2004 Elsevier Inc. All rights reserved.

Keywords: Hopf bifurcation; Periodic solution; Navier–Stokes equation; Bessel-potential space

1. Introduction

Bifurcation analysis plays a crucial role in understanding qualitative changes of flow regimes of oceanic and atmospheric circulation equations. From the view point of numer-

Corresponding author. *E-mail address:* zhimin@soton.ac.uk (Z.-M. Chen).

 $0022-247X/\$$ – see front matter \odot 2004 Elsevier Inc. All rights reserved. doi:10.1016/j.jmaa.2004.09.062

ical experiments, bifurcation phenomena relating to ocean circulation were examined by Charney and DeVore [2], Veronis [19,20], and Pedlosky [15] to determine the occurrence of multi steady-state solutions, whereas Jiang et al. [9], Jin and Ghil [10] predicted periodic circulatory motions.

Because wind-driven circulation problems embody extremely complex physical mechanisms, to aid understanding it is beneficial to study simplified models (see Ghil and Childress [6], Lions et al. [13,14], and Pedlosky [16]). For example, Pedlosky [16] developed a suitable mathematical model describing mid-latitude wind-driven circulation through a quasi-geostrophic approximation. This model describes the dynamics of circulation flow-driven by a wind stress and influenced by Ekman friction layers, bottom topography, and a *β*-plane approximation of the Coriolis force.

The purpose of this study is to derive a general Hopf bifurcation theorem of the twodimensional simplified wind-driven circulation equation (see Pedlosky [16, Eq. (5.2.22)]) which can be expressed in the following dimensionless vorticity formulation:

$$
\partial_t \Delta \psi + r \Delta \psi - \frac{1}{R} \Delta^2 \psi + \beta \partial_x \psi + J(\psi, \Delta \psi + \eta_B) = \beta \operatorname{curl} \tau \quad \text{in } \Omega \tag{1}
$$

together with the free slip boundary condition

$$
\psi = 0, \quad \Delta \psi = 0 \quad \text{on } \partial \Omega. \tag{2}
$$

Here $J(\psi, \phi) = \partial_x \psi \partial_y \phi - \partial_y \psi \partial_x \phi$ is the advection operator, $\psi = \psi(x, y, t)$ describes a geostrophic stream function, $\tau = (\tau_1(x, y), \tau_2(x, y))$ is a steady-state wind stress applied on the circulation basin *Ω*. The *x*-axis points eastwards and the *y*-axis northwards. The zonal and meridional velocity components *u* and *v* are given by

$$
u=-\partial_y\psi,\qquad v=\partial_x\psi,
$$

and the relative vorticity ζ is defined by $\zeta = \Delta \psi$.

The mathematical formulation (1) is a quasi-geostrophic approximation of the shallow water equation under the effect of a Coriolis force (see [16]). The parameter $\eta_B = \eta_B(x, y)$ is a function measuring the topography height of the bottom of the original fluid domain. Control parameters defining the fluid motion are Reynolds number *R*, the Ekman number *r* measuring the effects of friction arising from the top and bottom Ekman layers of the fluid, and the number *β* is derived from a *β*-plane approximation of the Coriolis force. The interested reader may find further descriptions of this equation and the theoretical background by consulting Pedlosky [16, pp. 260–261].

The existence of bifurcating periodic solutions branching off a basic steady-state solution of an ordinary differential equation system was first derived by Hopf [7,8]. The Hopf bifurcation theorem of Navier–Stokes equations is well-established under the assumption that the nonreal eigenvalue of the linearized spectral problem satisfies the simplicity condition and a condition relating to the transversal crossing of the imaginary axis at a critical value of Reynolds number (see Yudovich [22], Joseph and Sattinger [11]). Equation (1) is similar in form to the vorticity formulation of Navier–Stokes equations. Thus the Hopf bifurcation theorem derived in [11] is applicable to (1) to ensure the existence of bifurcating periodic solutions as the Reynolds number varies across a critical value *Rc* for constant Ekman number. However (1) also depends on the Ekman number *r* and this introduces

Download English Version:

<https://daneshyari.com/en/article/9503247>

Download Persian Version:

<https://daneshyari.com/article/9503247>

[Daneshyari.com](https://daneshyari.com)